

Original Research Article

Geophysical Evaluation of Central Nigeria Earth Tremor Activities Using High Resolution Airborne Magnetic Data

ABSTRACT

The aim of this study is to investigate the actual causes of earth tremors in Central Nigeria using geophysical technique. Since 1933 till date (over 88 years), small to moderate magnitude earthquakes have either been observed or instrumentally recorded in Nigeria. However, the specific causes of the intra plate earthquakes in Nigeria is still not well known. In this study therefore, a high-resolution aeromagnetic data analysis was performed to evaluate the various theories proposed as the causes of tremor activity in central Nigeria with a view to ascertaining the actual causes of earthquakes in the West African country that lies on a stable crust. The total magnetic intensity data acquired from the Nigerian Geological Survey Agency (NGSA) in the Central Basement complex of Nigeria, was processed and filtered using the tilt derivative, lineament analysis, and advanced Analytic Signal- Hilbert Solution filter. To define the trends of both major and minor structure, a lineament plot was extracted and plotted. The analysis revealed a dominant NE-SW and NW-SE trending structure that is distributed throughout the study area. The regional North Atlantic Romanche fault system, which extends from the offshore (Niger Delta) into the continental crust, was mapped and is thought to extend from Gwagwalada in the study area's south, through Gwarinpa and Mpape, to Kwoi, and beyond Kafanchan. Findings from our study have indicated that, areas prone to Earth tremors are linked to the Romanche fault by numerous NE-SW and NW-SE faults that act as conduits for seismic energy that causes tremor. Using lineament density analysis, tremor active areas have been interpreted to be highly deformed; thus, seismic energy induced into the environment would result in tremors. This study demonstrates that the regional North Atlantic Romanche fault system is the primary source of seismic energy that causes the earth tremors in central Nigeria.

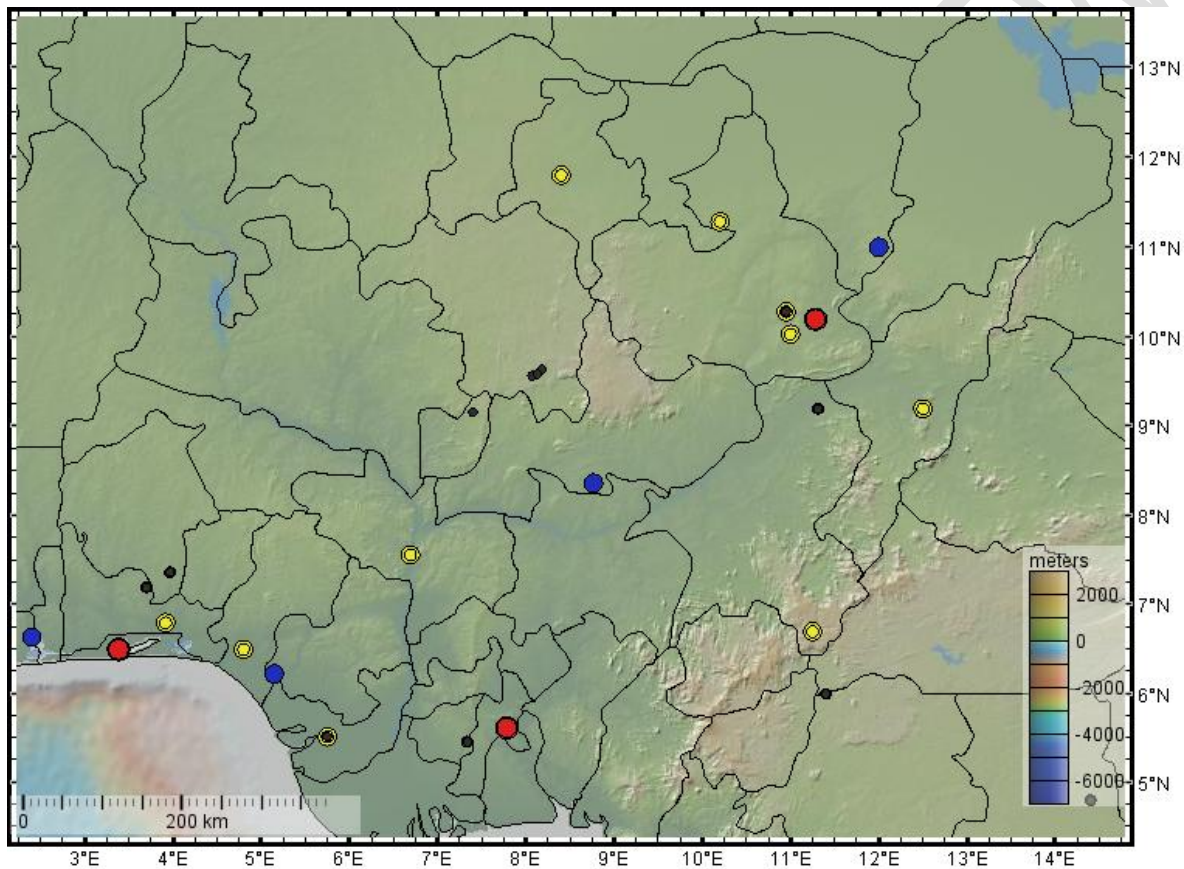
Keywords: Aeromagnetic, Central Nigeria, First Vertical Derivative, Lineament, Causes of Earth Tremors.

1. INTRODUCTION

Earth tremors are minor involuntary seismic movement of the earth crust. They are often described as slight earthquake with lesser magnitude measuring from Magnitudes 1 – 4.5. An earthquake could be generally defined as a sudden vibration, sometimes violent, of the earth's surface that follows a release of elastic energy in the Earth's crust [1]. In areas felt, it is described as a rapid rattling of the earth leading to falling of objects from shelves, breaking of glasses, vibration of the crust and uprooting of trees.

Seismic activities (tremors) have been reported in different parts of Nigeria since 1933 [2]. Previous investigations include those of [2], [3], [4], [5], [6], [7], [8], [9], [10], [11], [12], [13], [14], [15], [16], [17], [18], [19], [20], [21] and [22]. Most of these works were based on the regional studies of seismic activities in Nigeria except for the work of [21], who studied the effect of tremor event in Kwoi area in 2016. The earliest tremor in Nigeria was recorded in 1933 Warri (Delta State) and Ohafia (Abia State)

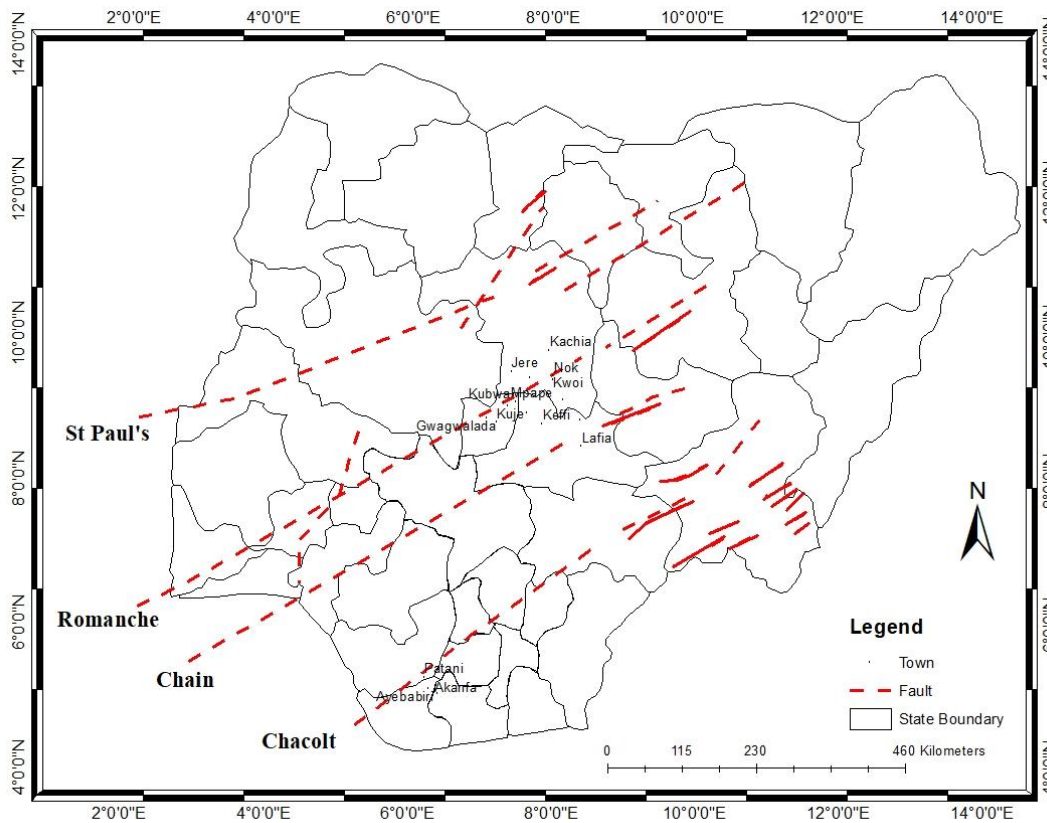
41 while the most recent was felt within Abia State, Nigeria in 2020. [4] stated that Nigeria sits on a stable
 42 Precambrian Basement complex rocks hence, the region is not expected to experience earth tremors.
 43 However, the country has experienced tremor with magnitudes ranging from 2.9 to 4.5. [6] ; [7] ; [8] and
 44 [5] reported that Nigeria has experienced about 31 earth tremors activities between 1933 and 2015. [20]
 45 stated that ten seismic activities were experienced in 2016, in addition to the earth tremor of Abuja in
 46 2018 and 2019; and Abia State in 2020; making a total of about 44 notable seismic activities in Nigeria.
 47 The Centre of Geodesy and Geodynamics, reports that the average value of intensities for Nigerian
 48 Earthquakes ranges from III –VI on the Modified Mercalli Scale, with the recorded Mpape earth tremor a
 49 magnitude of 3.1. The earth tremors of 2018 at Mpape, Federal Capital Territory and those of 2016
 50 whose epicenters were located at Kwoi in Kaduna State respectively, had moderate impacts on
 51 structures. Figure 1 shows the epicenters of some of the selected earth tremors that occurred in Nigeria
 52 between 1933 and 2021.
 53



54
 55 **Figure 1: Epicenters of earthquakes in Nigeria from 1933 to 2021. Legend: Black dots= $3 < M < 3.9$;**
 56 **Blue dots= $4 < M < 4.9$; Yellow dots= $5 < M < 5.9$; Red dots= $6 < M < 6.9$**

57 There are numerous schools of thought on the likely causes of the seismic activities in Nigeria. [22] , [11] ,
 58 [13] ; [9] ; [10] accredited the causes of earth tremors in the country to earth movement associated with
 59 NE-SW trending fractures and zone of weakness extending from Atlantic Ocean into Nigeria. This Atlantic
 60 transform fractures include the St Paul, Romanche, Charcot and Chain fractures zones (figure 2). The
 61 presence of these fractures zones which prominently traverse the Western half of Nigeria has been
 62 proposed by [11] ; [12] and [7] as the source of seismic activities experienced in the western half of
 63 Nigeria. The findings of [22] study on earth tremors that occurred at Kwoi are consistent with the theory
 64 of tectonic origin. This is because all of the epicenters plotted along an extrapolated trend coincided with
 65 the North Atlantic regional fracture zone.

66 Another school of thought by Onuoha (1988), attributed the earth tremors to partial reactivation of fossil
 67 plate boundaries. [20] also talked about the theory of stress transfer from plate boundaries. [5] and [8]
 68 correspondingly, argued that stresses that build up around plate boundaries could travel towards the
 69 center of the plates triggering intraplate tremors especially in pre-existing faults.
 70



71

72 **Figure 2: North Atlantic Fault Systems and their Extensions into Nigeria's Landmass**

73 [20], [2]; [8]; [5] agreed that tremors activities resulted from regional stresses in the crust since Nigerian
 74 Basement Complex rocks occur between the West African Craton and Congo Craton. [23] inferred a
 75 principal stress acting WNW-ESE in the basement rocks and they accredited this stress as the major
 76 cause of the tremor. Their thought is in agreement with the regional stress models of [24]; [25]; [26] and
 77 [27]. [20], [2]; [8] also proposed the theory of zone of weakness stating that when a magmatic intrusion
 78 is present in stable rock, the difference in geophysical properties can cause localized stress
 79 concentrations particularly when the intrusion is weaker than the surrounding rock. They stated that
 80 Initially, the mafic intrusion is more rigid than the surrounding rock but over time, the intrusion can
 81 become much weaker than the surrounding rock, hence leading to breaking that resulting is the release to
 82 seismic waves.

83 [15] stated that evidence has shown that anthropogenic activities such as oil and gas exploration, could
 84 trigger earthquake and Nigeria is among the top oil producing countries. More localized study of the
 85 tremor affected areas by Nigerian Geological Survey Agency and Centre of Geodesy and Geodynamics,
 86 Nigeria stated that the Mpape and Kwoi tremor activities could have resulted from human activities such
 87 as mining (blasting processes) and borehole (water withdrawal) and resettling of geologic formations.

88 As there is no confirmatory existing study with definite results about the causes of the intraplate earth
 89 tremors in the Central part of Nigeria, this study therefore assesses three major schools of thought: the
 90 theory that attributes tremor activity to a NE-SW trending fracture extending from the North Atlantic Ocean
 91 into Nigeria, the theory of zone of weakness, and the theory of human activities. As the airborne magnetic
 92 data has a rapid coverage of large area and cheaper to acquire, hence, the method was used in our

93 study. The aim of this study therefore, is to investigate earth tremor activities in the Central Nigeria using
 94 geophysical tool, in order to ascertain their actual causes.

95 2. THE STUDY AREA

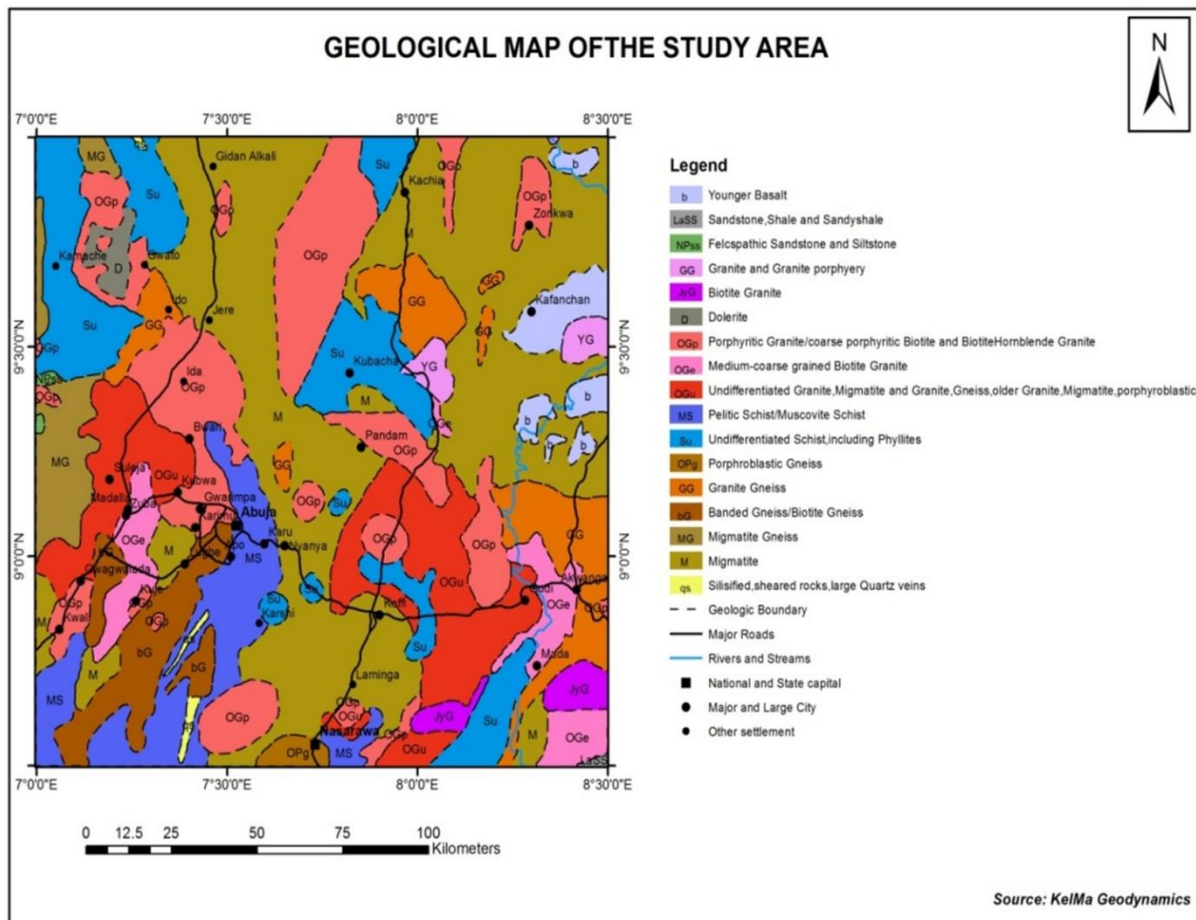
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97 The study area is located in the North Central part of Nigeria. It is bounded by Latitudes 7.00 °N and 8.30
 98 °N and Longitudes 8.30 °E and 10.00 °E. it covers a total surface area of about 27,225 km². Four major
 99 states, including Federal Capital Territory, Nassarawa, Kaduna and Niger State (Figure 3) is covered in
 100 this study. The two major communities of interest are Mpape (Federal Capital Territory) and Kwoi
 101 (Kaduna State).

102 The study area falls within the Basement complex of Nigeria, it is part of an upper proterozoic to lower
 103 phanerozoic mobile belt situated between West African and congo **craton**. The belt is believed to have
 104 evolved by plate tectonic process resulted from continental collision between the passive continental
 105 margin of the Tuaraep Shield [28] ; [29] ; [30] . The rocks of the basement complex have been subdivided
 106 by [31] , [32] ; [33] , [34], [35]; Geologically, the study area is underlain by metasedimentary rocks
 107 (migmatite, gneiss, schist and quartzites) and these metasedimentary rocks have been intruded by
 108 granitic rocks (Older and Younger Granite Suite). The rock types constitute the major Basement Complex
 109 rocks as defined by [31] , [32] ; [33] , [34], [35];. The granite rocks are well exposed in outcrop, while the
 110 metasedimentary rocks are majorly exposed at road cuts and quarries.

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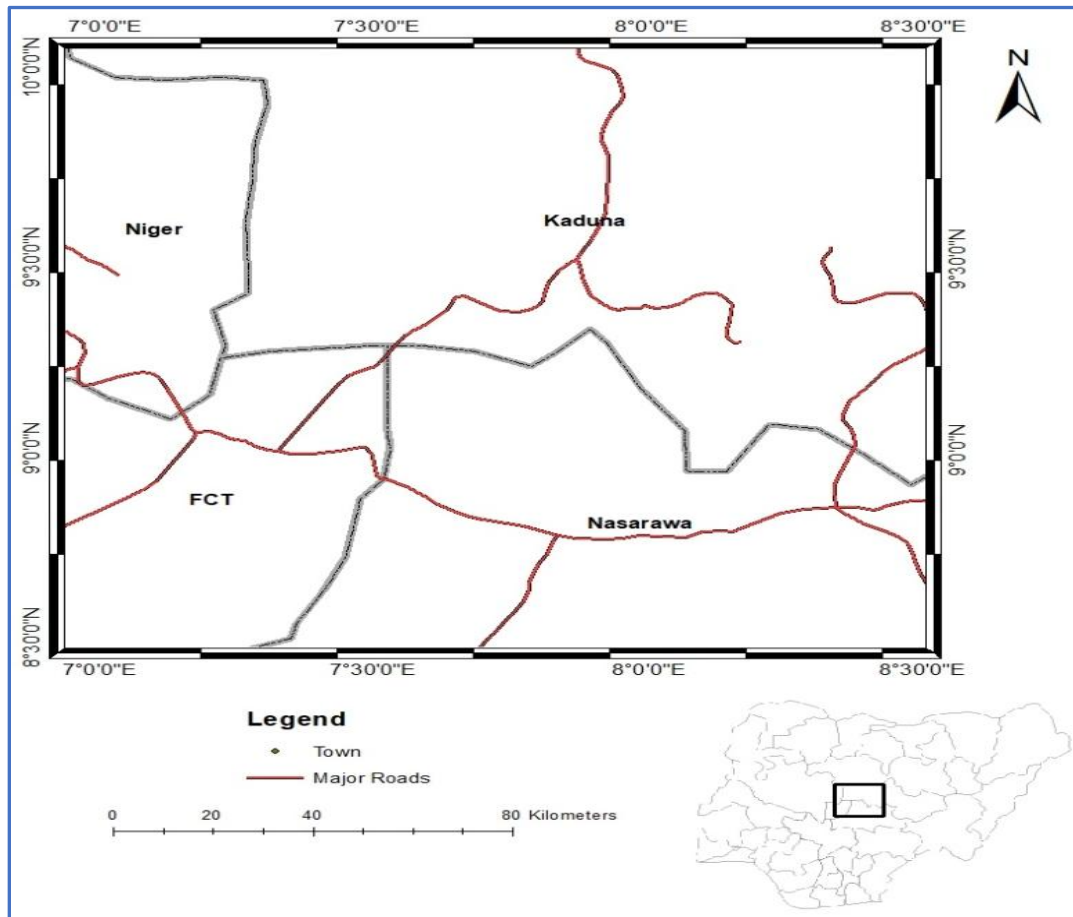


113

114 **Figure 3: Geology Map of the Study Area**

115

116 Figure 4 is the accessibility map, it explained how the study area can be accessed through major roads,
 117 minor roads and footpaths during geologic mapping.



118

119 **Figure 4: Figure 4: Accessibility map of the Study Area**

120

121 3. MATERIALS, METHOD AND DATA PROCESSING

122 The airborne magnetic dataset used for this study was acquired by Fugro Airborne Survey and Nigerian
 123 Geological Survey Agency. The data was acquired along a NW – SE flight lines at 500 m line spacing and
 124 80 m terrain clearance. The climate and average weather within the study area are of two seasons: from
 125 April to October is characterized with a low temperature and the wettest month of June; while the dry
 126 season runs from November to March with mid-day temperature that rise above 38°C (100°F) but relatively
 127 cool night dropping as low as 12°C (54°F). The data were acquired under these weather conditions.

128

129 The airborne magnetic dataset consists of nine (9) index sheets including Bisnini-165, Kachia-166,
 130 Kafanchan-167, Abuja-186, Gitata-187, Jemaa-188, Kuje-207, Keffi-208 and Akwanga-209. Each data
 131 sheet covers an estimated 3,025-kilometer square. The total magnetic intensity grid was developed by
 132 employing a minimum curvature algorithm at 100 m grid cell size. The digitized data were filtered by
 133 utilising a low pass Fourier domain sub-routine filter to exterminate undesired wavelengths and to pass
 134 longer wavelengths. The aeromagnetic dataset was subjected to Reduction-To-Pole (RTP) transformation
 135 to reduce appreciably the polarity effects. A computer program, Geosoft (Oasis Montaj, version 8.4), was
 136 utilised to generate the residual magnetic anomaly values. This was achieved by subtracting regional
 137 fields values, from the measured intensity of the total field, at the grid cross point.

138 Tilt derivatives was applied to enhance structural interpretation. It is extremely useful in mapping out
 139 shallow basement structures. Tilt derivative was introduced by [36] and later modified by [37], its defined
 140 as:

$$141 \quad TDR = \tan^{-1} \frac{\frac{\partial T}{\partial z}}{\sqrt{\left[\frac{\partial T}{\partial x}\right]^2 + \left[\frac{\partial T}{\partial y}\right]^2}} \quad (1)$$

142
 143 Where; T is the observed magnetic anomaly at (x, y) .
 144 Qualitative interpretation of magnetic anomaly attempts to give a reliable estimate of the direction, size,
 145 shape and depth of magnetic sources. It is a complementary to qualitative interpretation. This technique
 146 is broadly characterized as curve matching, forward or inverse modeling [38]. In 1982, Thompson
 147 developed a technique known as Euler deconvolution method, to process magnetic data and to convolve
 148 them to a point source at depth. Equation 2 is known as Euler deconvolution equation. For an arbitrary
 149 specific structural index the depth and location for various targets can be resolved from these first order
 150 derivatives (x, y and z).
 151

$$152 \quad (x - x_0) \frac{dT}{dx} + (y - y_0) \frac{dT}{dy} + (z - z_0) \frac{dT}{dz} = N(B - T) \quad (2)$$

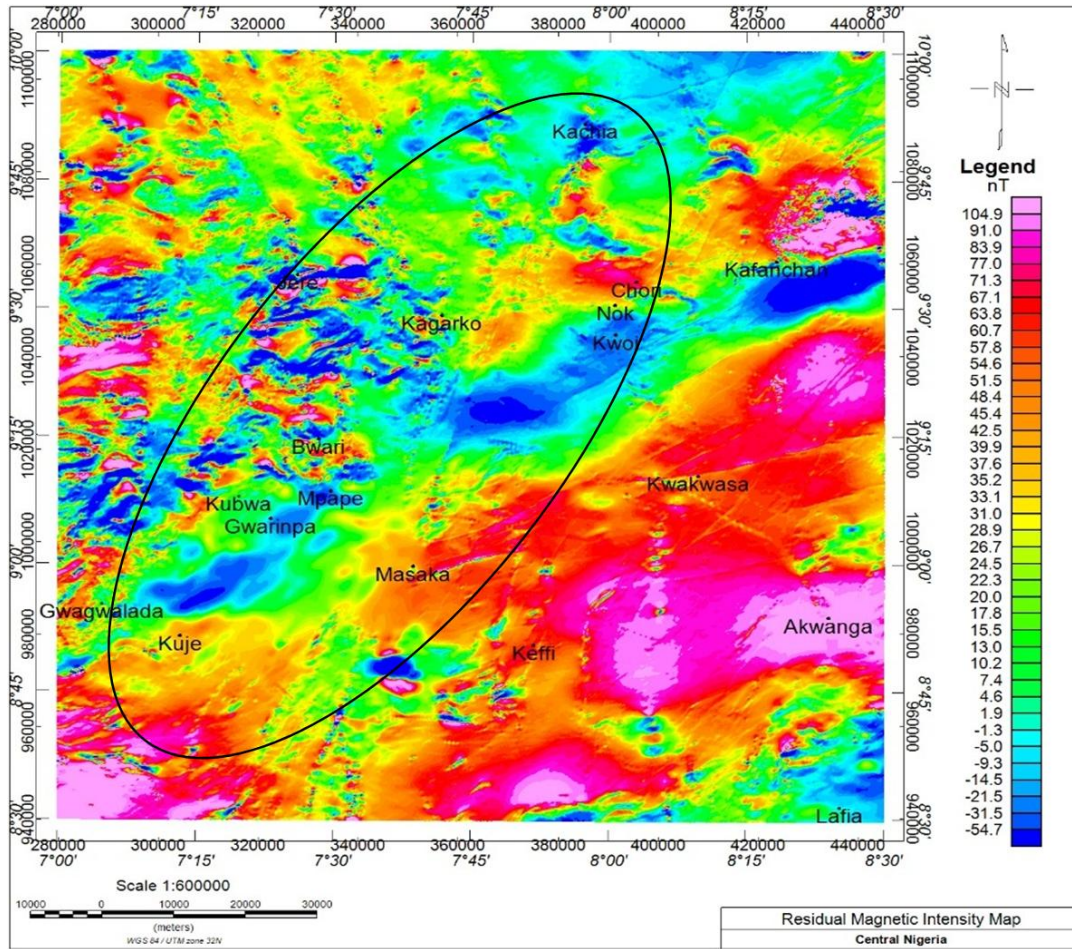
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 154 Where; (x_0, y_0, z_0) is the source of magnetic anomaly (x, y, z) is the total magnetic field detection points, B is
 155 the regional magnetic field. N is the fall-off rate magnetic field measurement else known as the structural
 156 index (SI).
 157

158 4. RESULTS AND DISCUSSION

159 4.1 Theories related to NE-SW Atlantic faults and Human Activities

160
 161 Energy are transmitted into smaller fault lines in NE direction base on the angler relationship of
 162 Romanche fault and the localized smaller fault lines. The residual magnetic intensity map (figure 5)
 163 reflects the local magnetic anomaly within the study area. The residual magnetic anomaly map was
 164 computed by applying IGRF to the acquired total magnetic intensity dataset and polynomial fitting applied
 165 to the IGRF result. The residual magnetic intensity has a range of -54.7 nT to >104.9 nT,, the high
 166 magnetic intensity is observed at the southeastern part of the study area (Akwanga, Kwakwasa, Masaka
 167 and Keffi). The northern and central part of the study area is characterized by intermediate to low
 168 magnetic intensity (around Kwoi, Nok, Kafanchan, Kachia, Jere, Bwari, Mpape, Gwarinpa). The
 169 northwestern part of the study area (around Jere, Bwari, Kubwa area) is characterized by bipolar
 170 magnetic bodies, indicative of intrusive rocks (granites).
 171

172 Numerous extensive faults are observed within the study area. The regional Romanche fault is
 173 interpreted within the study area; it cuts into the study area at Gwagwalada, passing through Gwarinpa,
 174 Mpape, Kwoi, Nok and exist the study area through Kafanchan. The Romanche fault has a NE-SW trend
 175 and extends from the North Atlantic Ocean. Minor faults are also observed at the southeastern part of the
 176 study area (Akwanga, Kwakwasa, Masaka and Keffi), with a dominant NE-SW trend and minor NW-SE
 177 trends. Numerous low magnitude faults are also observed around Kachia. The intensity of the bipolar
 178 structures within Jere, Bwari, Mpape, Kubwa and Kagarko is indicative of intense deformation and
 179 probable weakness of subsurface geology.



180

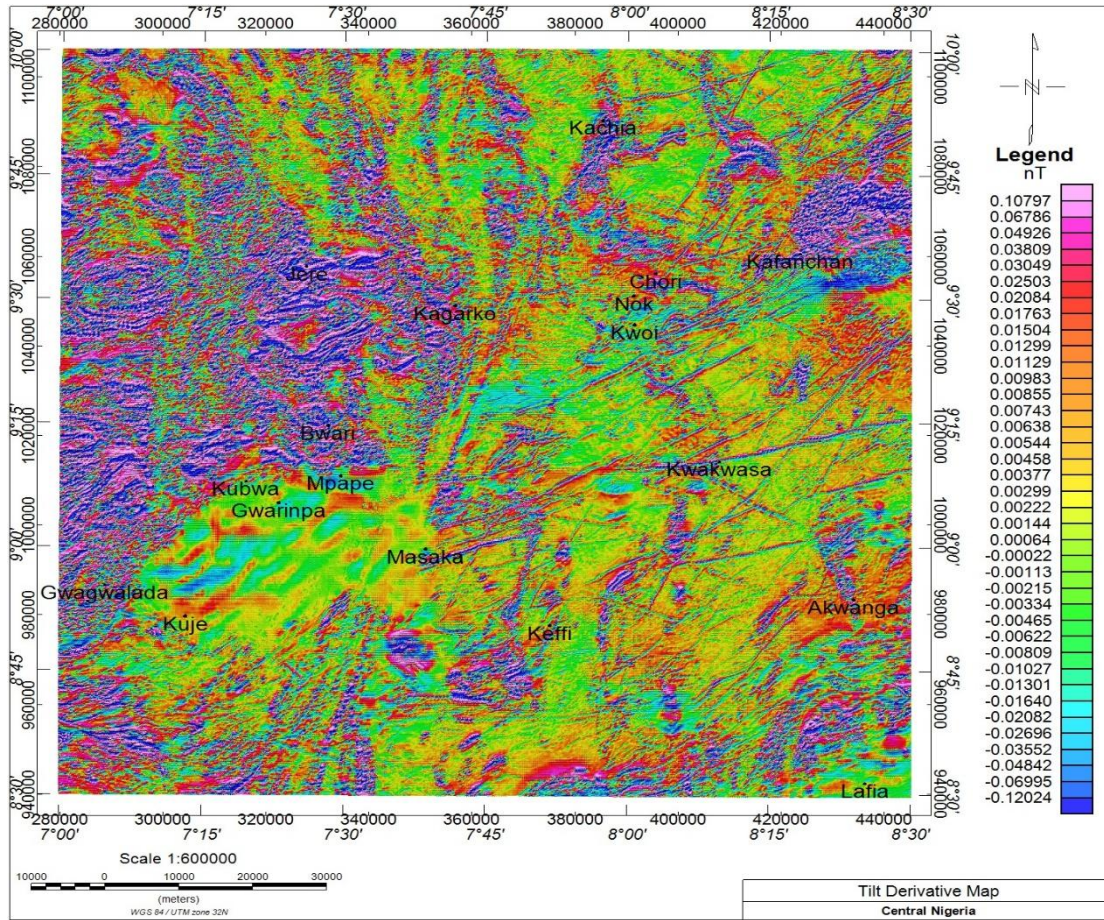
181 **Figure 5: Residual Magnetic Map of the Study Area Showing Interpreted Fault. The major region is**
 182 **enclosed under the black ellipsoid**

183

184

185 Tilt derivative (TDR) enhance and sharpen up anomalies over causative bodies and tend to reduce
 186 complexity, allowing a clearer imaging of the causative structures. The TDR map (Figure 6) shows
 187 inferred faults, fractures, folds, contacts and to some extent the shape of some lithologic contacts which
 188 indicate structural features. The tilt derivative map better enhances the faults within the study area, the
 189 fault has a regional dominant NE-SW, with minor NW-SE trend. The NE-SW trending structures is
 190 interpreted as the older structural trends while the NW-SE is the younger; this is based on their cross-
 191 cutting relationship as observed around Kwoi, Nok, Chori, Kafanchan, Kwakwa, Keffi, Masaka and
 192 Akwanga.

193



194

195 **Figure 6: Tilt derivative (TDR)Map of the Study Area Showing Interpreted Fault**

196 Qualitative interpretation of both the residual and tilt derivative map defied numerous major faults within
 197 the study area, although the tilt derivative map better enhanced the structures.

198
 199
 200 The regional $F1$ fault ($F1$ is the regional Romanche fault in figure 7) is interpreted to have demarcated the
 201 study area into two different structural complexes with the eastern structural complex is characterized by
 202 extensive faults lines, while the western portion is characterized by bipolar structures separated by NNW-
 203 SSE faults. The regional fault $F1$ is interpreted is better visualized in the tilt derivative map than the
 204 residual magnetic map, the fault extends from the south of Masaka, through Kagarko to Kachia. Parallel
 205 faults trending NE-SW is observed at the eastern part of the study area, extending from Kwoi, Nok, Chori
 206 to Kafanchan, these faults looks to be terminated by the regional $F1$ at the central part of the study area.
 207 The study area is also divided into the northern structural complex and the southern structural complex by
 208 the Romanche Fault system. The northern structural complex is characterized by high occurrence of
 209 bipolar structures while the southern part is characterized by low occurrence of bipolar structures, this is
 210 indicative of different tectonic regimes and high intensity of tectonic activities in the past.

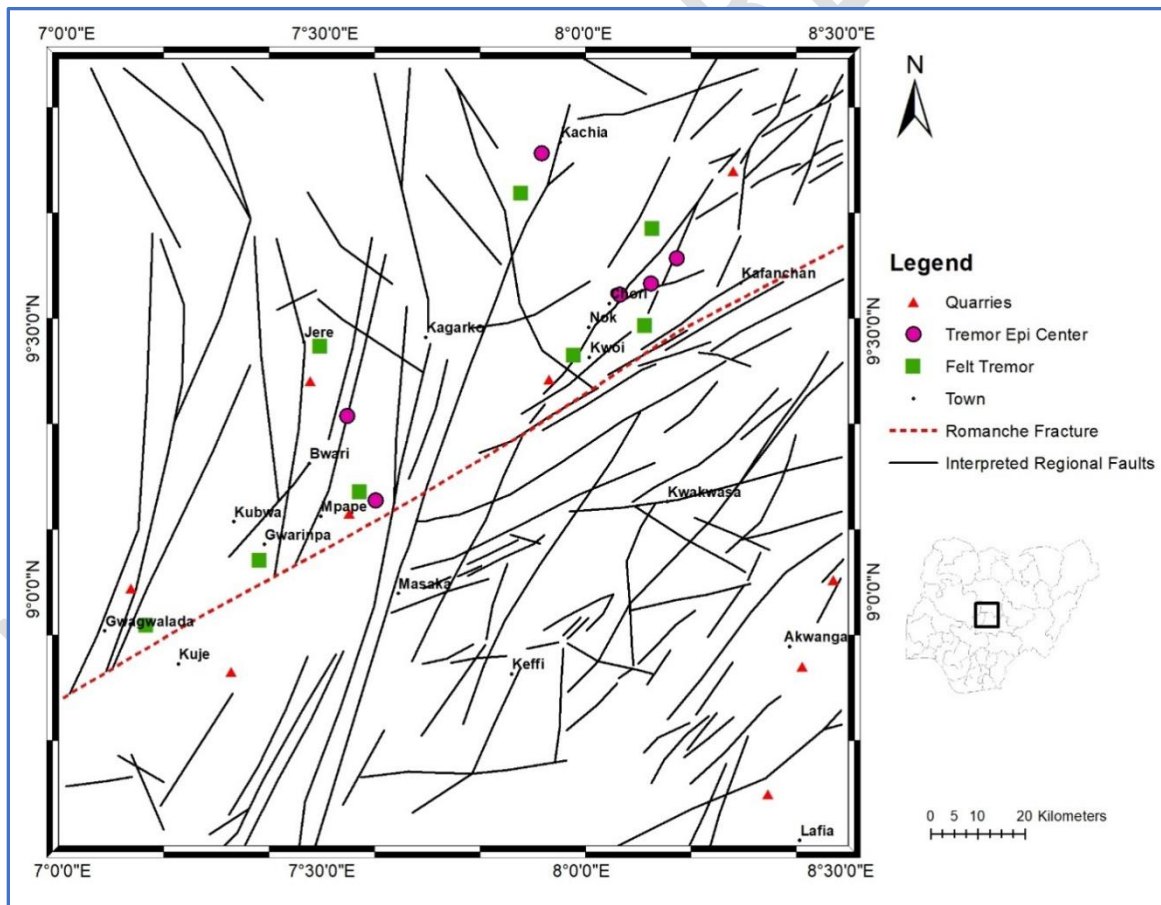
211
 212 The qualitative interpreted structures where used to produce a structural map of the study area (Figure 7).
 213 From the interpreted structural map, the oldest structure Romanche Fault, and the three deformation
 214 events were identified based on their structural relationship. The early deformation event (D1) is defined
 215 to be a ductile deformation producing folds whose axial planes trends NNE-SSW, the second deformation
 216 event resulted in the reactivation of the D2 structures particularly the folds in the NE-SW direction. The
 217 third deformation event produced a transgressive deformation starting with ductile deformation resulting in
 218 the development of NE-SW to EW oriented structures, brittle deformation occurring at later stages of the

219 D3 event producing fractures oriented in the NW-SE and EW direction which are more associated with
 220 the reactivation of D1 and D2 faults and fractures.

221

222 The epicenters of the earth tremors as located by the local seismic stations operated by the Centre for
 223 Geodesy and Geodynamics, Toro, Bauchi State, Nigeria, where plotted on the interpreted structural map
 224 so as to understand the structural relationship to the seismic activities. It was observed that all the earth
 225 tremor epicenters are located around the regional NW-SW structural joints which in turn are all connected
 226 to the continental Romanche Fault line. These therefore, indicate that, the transfer of energy from the
 227 continental Romanche Fault line into localized regional fault lines could be a major cause of the tremor
 228 activities. This is evident around Mpape, Kwoi and Kachia areas, where the epicenters sit close to
 229 structural joints of faults which are all connected to the continental Romanche Fault line (Figure 7).
 230 Gwagwalada and Jere areas where earth tremors have also been felt, are seen to be located around a
 231 regional fault joint that is also connected to the continental Romanche Fault line. It was also observed that
 232 seismic activities are concentrated at the northern part of the Romanche fault line, with no record of these
 233 activities in the south. This can simply be explained by employing the angular relationships between the
 234 continental Romanche Fault line, the localized regional faults and, the direction of the transferred seismic
 235 energy. When the continental Romanche Fault line comes into contact with regional low angular NE-SW
 236 trending fault lines at its structural joints, the energy is transferred to a smaller fault line in the NE
 237 direction based on the angular relationship of the faults. Also, small seismic energy transmitted through
 238 the continental Romanche Fault line would be larger when transmitted through the localized regional fault
 239 lines; hence, resulting in seismic activities when it collides with the structural joints during the transfer of
 240 seismic energy.

241



242

243 **Figure 7: Interpreted Structural Map of the Study Area**

244 Quarry locations were also plotted on the structural map, to analyze the relationship between blasting
245 activities during mining and the tremors experienced within Central Nigeria. Numerous quarries were
246 located around the epicenters where earth tremors were felt as well as at the southern part of Romanche
247 fault. Some of these quarries are located close to regional fault lines, as observed around Jere, Mpape,
248 Kwoi and Gwagwalada; therefore, blasting activities could lead to the accumulation of stress within these
249 faults, which possibly will lead to seismic activities.

250
251 This theory anyway does not explain why earth tremors are not felt in the southern part of the study area
252 which also has numerous quarries located around regional fault lines. Therefore, the theory of
253 mining/quarries activities as being the major cause of seismic activities in the study is unlikely. This
254 position could better be understood, when we reflect on the numerous blasting activities dating back to
255 the mid-1990s that were never associated with any tremor activities within these areas. Thus, it is safe to
256 conclude that blasting activities during the mining process is an unlikely source of seismic stress that lead
257 to the earth tremors in Central region of Nigeria.

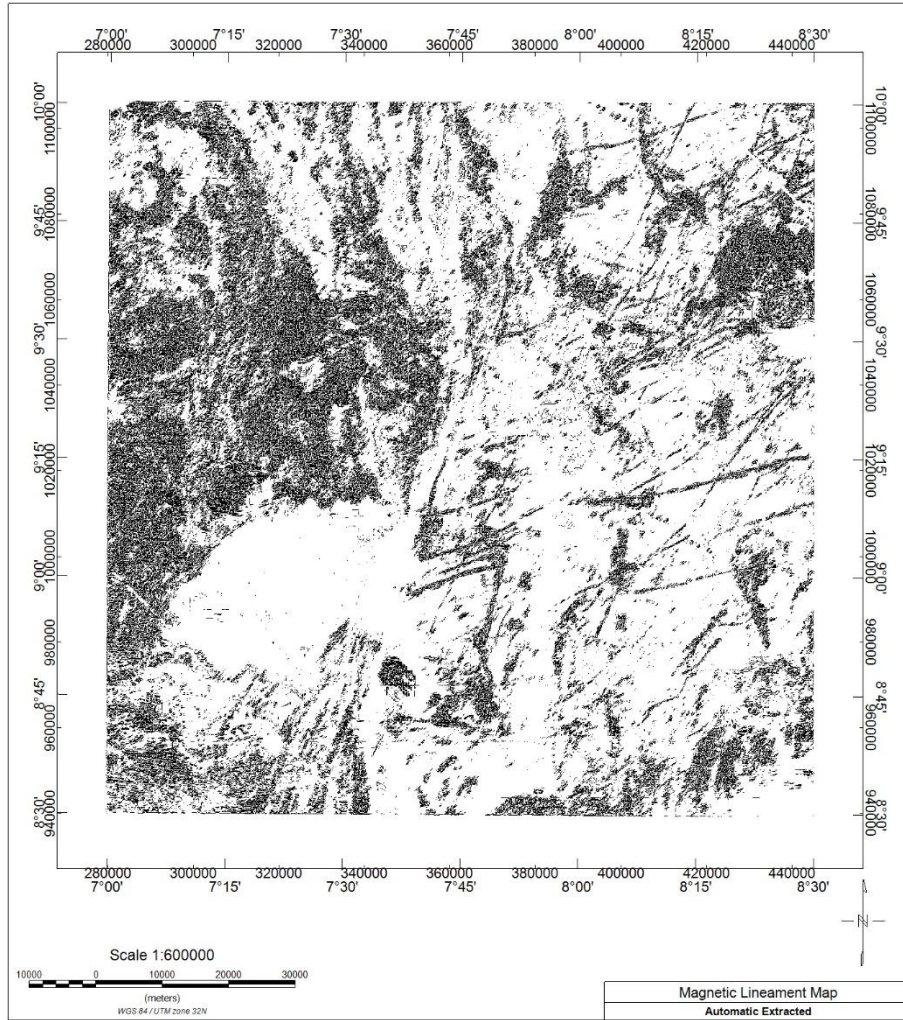
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261 **4.2 Theories related to zone of weakness**

262 Lineaments were extracted from the tilt derivative map to aid structural intensity to evaluate the
263 deformation/stability of the local geology. Lineaments are linear features on a landscape which expresses
264 the underlying geological structures such as faults, joints, and arched zones or even geological contacts
265 and could be of regional extent, usually in linear or curvilinear continuous or discontinuous over an entire
266 length [39].

267
268 The lineament map shows a dominate NE-SW and NW-SE structural trend, with minor E-W trends. The
269 extracted lineaments (figure 8) was used to compute the lineament density map (figure 9) of the study
270 area, which generally indicates the density of lineaments/structure at a particular location. The lineament
271 density is directly related to the structural complexity of an area, which is also directly proportional to the
272 stability of geologic structures. Areas with high lineament density indicate high deformation and low
273 structural stability, while in areas with less lineament density indicate low deformation and high structural
274 stability [40].

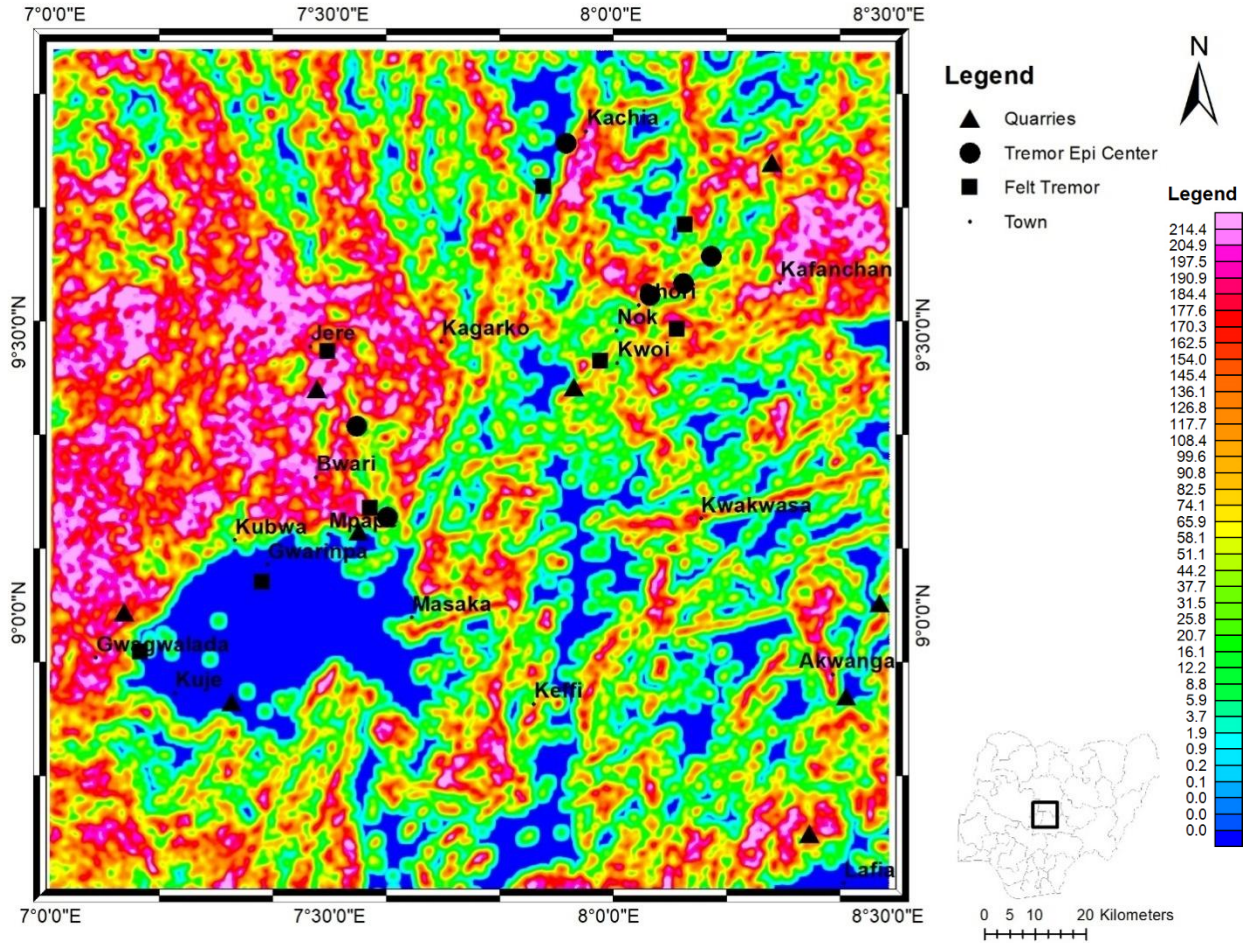


275

276 **Figure 8: Lineament Map of the Study Area**

277

278 Using figures 8 and 9, Jere, Bwari, Kubwa, Gwarinpa, Kagarko, Kwoi, Chori, Kafanchan, Gwagwalada
 279 and Kachia areas show high lineament density. Therefore, these areas are defined as low structural
 280 stability/ high zone of weakness regions. This is important because, based on their intense deformation,
 281 any seismic stress induced into the environment could result into earth tremors. Source of this seismic
 282 stress could be the continental Romanche Fault line or cumulative effect of blasting within the weakness
 283 zones.



284

285 **Figure 9: Lineament Density Map of the Study Area**

286 Areas around Akwanga, Masaka, Kwakwasa, Kuje, and Keffi also have notable zones of weakness (high
 287 lineament density regions), but not as much as the northern part of the study area. These findings also
 288 suggest that the zone of weakness within the local geology of the tremor active areas plays an important
 289 role in the earth tremor activities, and that the continental Romanche Fault line could be a major source of
 290 seismic energy, with localized regional fault lines serving as the primary conduit for this energy.

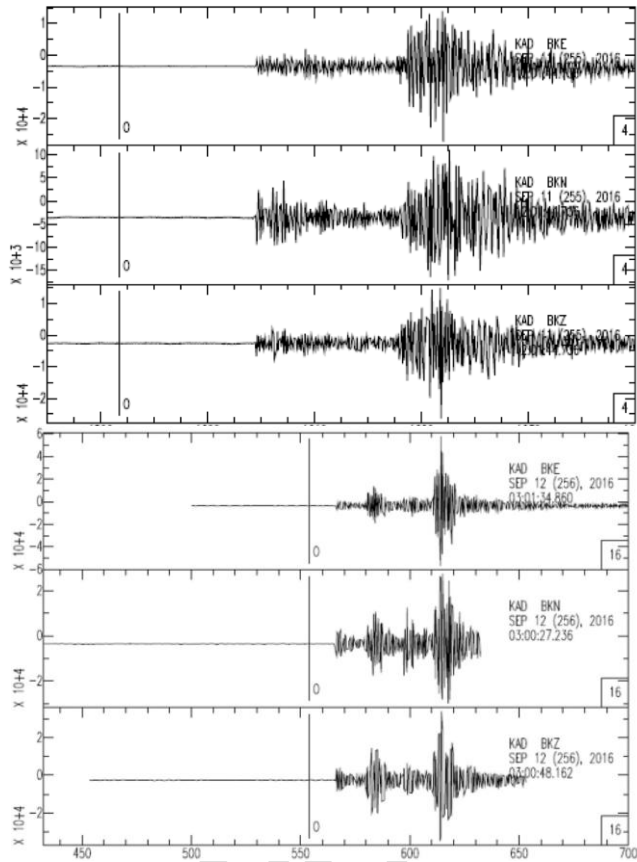
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292 **4.3 Depth to Epicenter of the tremors**

293

294 The focal depths of the 2016 Kwoi and 2018 Mpape earth tremor that were recorded by five seismic
 295 stations operated by the Centre for Geodesy and Geodynamics (CGG), Nigeria, range from 10-15km. The
 296 average macroseismic intensity value of III on the Modified Mercalli Scale was estimated from field work.
 297 The processed seismograms of the events by the staff of CGG, Toro, are shown in figures 9 and 10. The
 298 2018 Mpape events' parameters are presented in table 1.

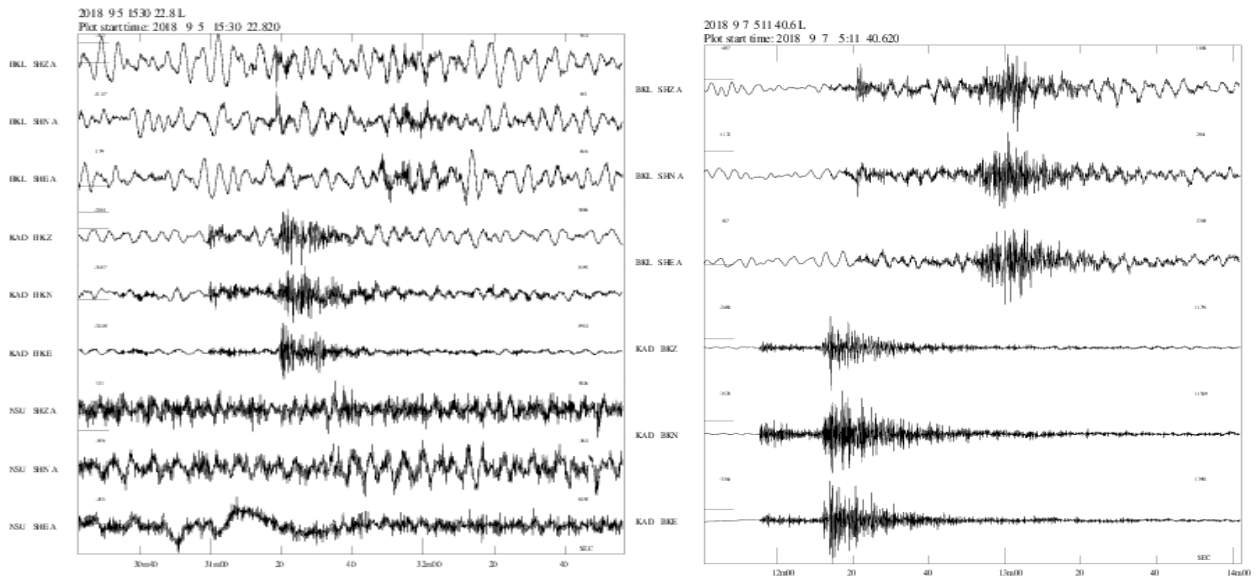
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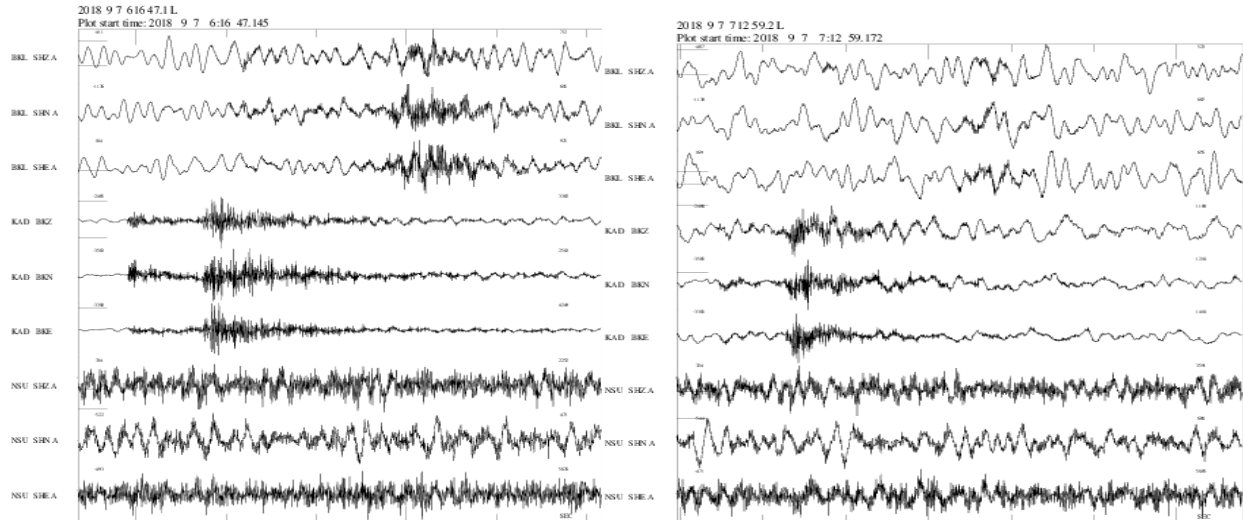
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301 **Figure 10: September 11 and 12, 2016 Kwoi tremor recorded in Nigeria (Source: Centre for Geodesy and**
 302 **Geodynamics, Toro)**

303



304



305
306 **Figure 11: (a) Seismogram of recorded foreshocks in Abuja on the 5th September, 2018(b)**
307 **Seismogram of Abuja Earthquake (main shock) on the 7th September, 2018 (c) Seismogram of**
308 **Abuja Earthquake (first aftershock) on the 7th September, 2018 (d) Seismogram of Abuja**
309 **Earthquake (second aftershock) on the 7th September, 2018 (Source: Centre for Geodesy and**
310 **Geodynamics, Toro)**

311 **Table 1: Parameters for locating the 2018 Mpape tremor events (Source: Centre for Geodesy and**
312 **Geodynamics, Toro)**
313

Event	Date	Time	Magnitude	Focal Depth (km)	Epicenter	
					Latitude (°N)	Longitude (°N)
Foreshock	05-09-2018	3:30	<2.5	-	-	-
Foreshock	05-09-2018	3:34	<2.5	-	-	-
Foreshock	05-09-2018	3:39	<2.5	-	-	-
Foreshock	05-09-2018	3:42	<2.5	-	-	-
Foreshock	05-09-2018	7:47	<2.5	-	-	-
Foreshock	05-09-2018	8:57	<2.5	-	-	-
Main Shock	07-09-2018	5:11	3.0	15	9.14	7.59
Aftershock	07-09-2018	6:16	2.6	12	9.03	7.50
Aftershock	07-09-2018	7:12	2.5	10	9.16	7.40
Main Shock	11-09-2016	2:28	2.8	10	9.57	9.64
Aftershock	12-09-2016	3:10	2.7	10	9.64	8.18
Aftershock	12-09-2016	3:11	2.9	10	9.59	8.13

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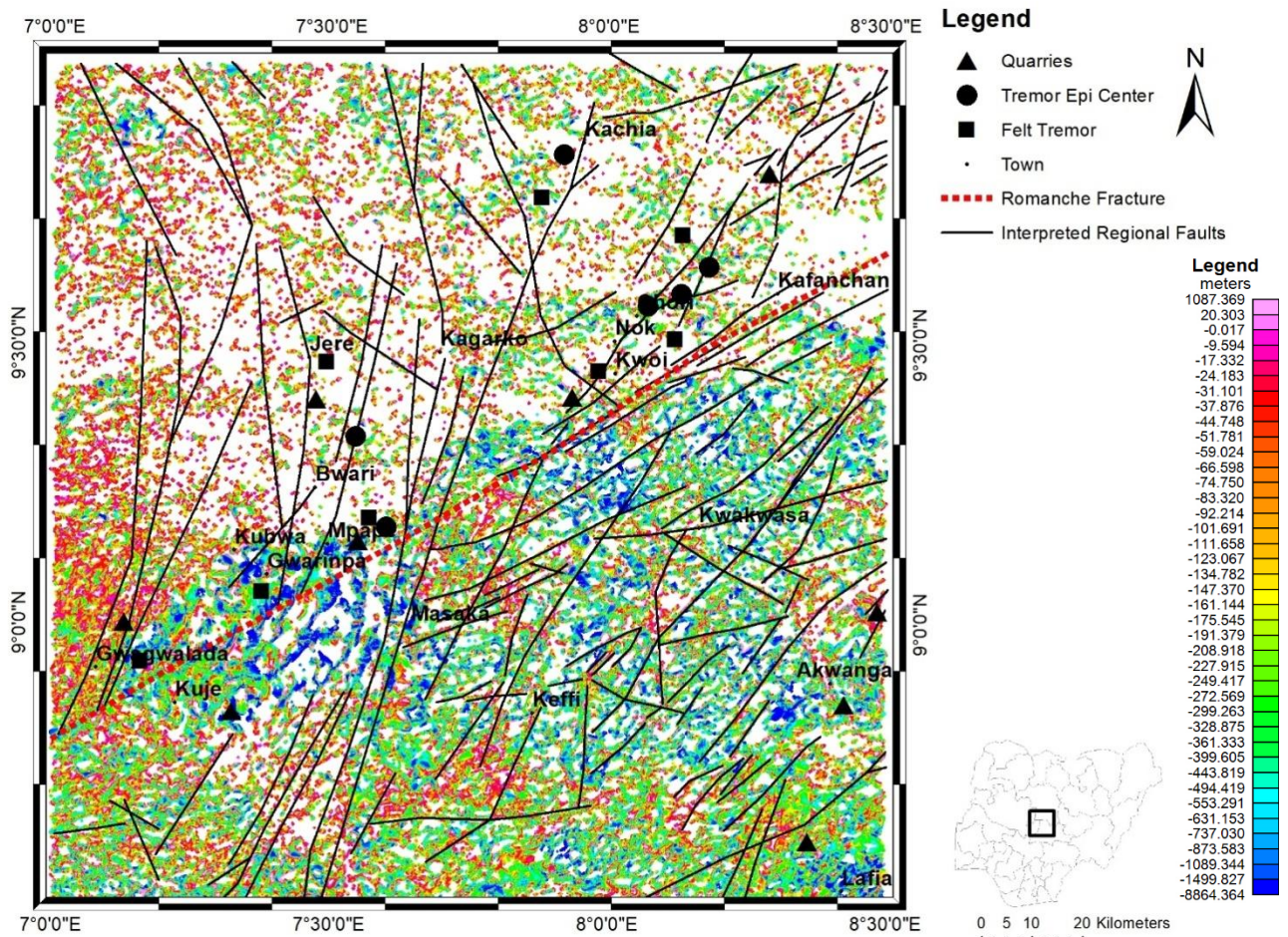
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316 The focal depths of the epicenters of the 2018 Mpape earth tremors range from 10 km to 15km. Depth
317 analysis was carried out using the magnetic dataset, by implementing the advanced Analytic Signal-
318 Hilbert Solution depth analysis to project the depth of interpreted regional fault structures around the
319 epicenters. This was done in order to evaluate the potential of the source of the seismic energy being
320 responsible for the observed earth tremors in the study area. The regional fault joint around Mpape area
321 was interpreted to have depths ranging from 8 km to 9 km (figure 12); these values are very close to the
322 values of the focal depths which range from 10 km to 12 km as estimated from the seismograms (table 1).

323 The faults joints around Kwoi, and Chori areas have a depth range of 7 km to 8 km, while the estimated
 324 focal depth using the seismograms was 10km.

325 It is therefore safe to conclude that, the relationship between the North Atlantic Romanche Fault and the
 326 continental regional faults system plays a key role in the source of seismic energy that resulted in the
 327 earth tremors in the central Nigeria.

328



329

330 **Figure 12: Depth to source map of the study area**

331

332

333 **5. CONCLUSION**

334 From the perspective of global seismicity scenario, Nigeria is located on a stable crust; however, from
 335 2016 to 2021, multiple earth tremors that were witnessed within the Central Nigerian region have elicited
 336 serious concern in the country. The average moment magnitude of the recorded and analyzed events is
 337 3.0 with average macroseismic intensity value of IV on the Modified Mercalli Scale. In an effort to
 338 ascertain the actual reason for the occurrence of the increasing seismic activities in the region of interest,

339 we evaluated the several theoretical schools of thought using qualitative and quantitative interpretation of
 340 airborne magnetic dataset.

341
 342 Findings from in-depth analysis have revealed that, the NE-SW trending fractures extending from the
 343 North Atlantic Ocean into Nigeria's landmass, and the zones of weakness, are the most likely sources of
 344 the seismic energy that resulted in the observed earth tremors in the central Nigeria. Findings from this
 345 study equally showed that the regional North Atlantic Romanche fault system is the primary source of
 346 seismic energy that causes the earth tremors in central Nigeria. It is therefore recommended that,
 347 relevant authorities take decisive measures to establish a robust and integrated monitoring scheme in the
 348 central Nigerian region, to help in better monitoring of seismic activities and their attendant seismic
 349 hazards.

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352 **COMPETING INTERESTS DISCLAIMER:**

353 Authors have declared that no competing interests exist. The products used for this research are
 354 commonly and predominantly use products in our area of research and country. There is absolutely no
 355 conflict of interest between the authors and producers of the products because we do not intend to use
 356 these products as an avenue for any litigation but for the advancement of knowledge. Also, the research
 357 was not funded by the producing company rather it was funded by personal efforts of the authors.

358

359 **AUTHORS' CONTRIBUTIONS**

360 Author A' designed the study, performed the statistical analysis, wrote the protocol, and wrote the first
 361 draft of the manuscript. 'Author B' and 'Author C' managed the analyses of the study. 'Authors A and C
 362 managed the literature searches, citations and references. All authors read and approved the final
 363 manuscript.

364 **ACKNOWLEDGEMENTS**

365

366 The authors sincerely appreciate and thank the Centre for Geodesy and Geodynamics, Toro, Bauchi
 367 State, and the Nigerian Geological Survey Agency, Abuja, both in Nigeria, for freely making available, the
 368 data used in this study.

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