

# **A Comparison of the Standardized Precipitation Index (SPI) and the Rainfall Anomaly Index (RAI) during the Meteorological Drought across Harohar-Punpun Basin (India).**

## **Abstract**

The term "drought" applies to a prolonged period when there is a water shortage because of insufficient precipitation, an excessive rate of evaporation, and excessive use of water from reservoirs and other storage, including ground water. It is a slow-onset phenomenon that can have significant social, economic, and environmental impacts. Both the frequency and severity of droughts are increasing globally due to manmade and natural factors. This study looked at how well SPIs (3, 6, 12-month time scales) and RAI (Rainfall Anomaly Index) performed at identifying drought occurrences over a period of 70 years (1951 to 2021) in India's Agro-Ecological Zones. The RAI and SPI values were computed using rainfall data from 9 meteorological stations located throughout the Harohar-Punpun Basin. According to the results, the RAI is more capable than SPIs in detecting historical records of actual occurrences. Additionally, RAI is more effective than SPIs for determining both short- and long-term droughts. Since RAI can better capture the true nature of the drought situation in the Harohar-Punpun Basin, it appears to be more responsive to drought circumstances. The primary causes of drought, according to the study, include minor variations in precipitation, willful ground water removal, changing cropping patterns, and substantial changes in land use. Regional planners and administrators will surely find use for this scientific study and integrated watershed management methods for sustainable water resource management, particularly for this region.

**Keywords:** Meteorological Drought, SPI (Standardized Precipitation Index), RAI (Rainfall Anomaly Index), Harohar-Punpun Basin.

## **Introduction**

Drought is indeed a prolonged period of dryness and water scarcity that affects various parts of the Earth's system, including the atmosphere, lithosphere, and hydrosphere. It is considered one of the most severe climatic events. Drought occurs when hydrological and meteorological conditions result in a deficit of water supply compared to the normal or expected conditions (Kim and Jehanzaib, 2020). arid areas are generally more prone to drought compared to other regions. This is because arid regions naturally have low levels of precipitation and limited water

availability. Unlike tropical wet and humid regions, where rainfall is relatively consistent throughout the year, arid areas depend on sporadic and often infrequent rainfall events to sustain their water supply (Sun et al., 2006). According to Bhunia et al. (2019), this extreme climatic event has typically been divided into four categories: hydrological drought, agricultural drought, meteorological drought, and socioeconomic drought. When the decrease in rainfall for a given time frame a day, a month, a season, or an entire year—falls below a predetermined threshold, which is typically outlined as a percentage of the long-term average, it is experiencing a meteorological drought. According to Selvaraju and Baas (2007), it usually refers to a long-term deviation from regular precipitation. According to the Indian Meteorological Department, meteorological drought occurs when rainfall over an area during the south-west monsoon season (June to September) is less than 25% of the average. For the purpose of carrying out proper management of water resources, a scientific research and study are urgently required to determine the drought occurrence rate, magnitude, probable return period, and recurrence interval of severe and intense drought episodes (Bhunia et al., 2019). According to the Intergovernmental Panel on Climate Change's 2007 report (Change, 2007), the average surface temperature of the Earth has risen by around 0.76°C during the previous century. By the 2080s, India's temperature is expected to climb by 2.7–4.3 °C, according to the IPCC's 2007 study. The group also predicted that by 2100, the sea level will rise by 88 cm and there would be a 6–8% increase in rainfall over India. According to the Irrigation Commission of India, a drought is any place where the annual rainfall is less than 75% of what is considered normal. According to relevance, duration, and timing of occurrence, there are generally three types of droughts.

Given that drought is linked to climatic occurrences, factors like rainfall, temperature, and stream flow can serve as reliable predictors of the presence or absence of drought (provide ref). Then, these indications can be transformed into drought indices that demonstrate the occurrence, extent, severity, and duration of a drought event (Zarger et al., 2011; Hayes, 2006). Both a single input variable and a mixture of hydrological variables can be used to create drought indices (Hao and Singh, 2015). The type of drought being studied and the problem to be addressed determines which hydrological factors should be used in the indicators in order to produce more certain conclusions. The region of interest and the availability of data also influence the choice of the drought index (Smakhtin and Schipper, 2008). There are a number of several drought indices available (Heim, 2002), but the Palmer Drought Severity Index (PDSI, Palmer, 1968), rainfall

~~anomaly index (RAI)~~, and the ~~Standardized Precipitation Index (SPI)~~, (Mckee et al. (1993) are now the most often used drought indices.

The SPI is multi-scalar because it combines cumulative precipitation deficiencies at different spatiotemporal scales (Hou et al., 2007). A key presumption in the calculation of the SPI is that droughts are primarily caused by rainfall variability, whereas other variables, such as temperature, are stationary and, as a result, do not fluctuate over time (Vicente- Serrano et al., 2010). The primary advantage of SPI over other indices is that it can identify drought at multiple ~~time~~ periods (1, 3, 6, 12, and 24 months), indicating that diverse types of droughts (meteorological, agricultural, and hydrological) may be tracked. However, according to Tirivarombo and Hughes (2011), the quality of a drought index output is only as good as the input data. The classification of the positive and negative severities of rainfall anomalies is done using the ~~Rainfall Anomaly Index (RAI)~~, which was created by Rooy (1965). Because it just requires precipitation data, it is regarded as an index of outstanding procedural simplicity (Freitas, 2005; Fernandes et al., 2009). Regarding the Keyantash and Dracup (2002) stated evaluation criteria for drought indices, the RAI offers a higher degree of transparency and tractability and requires a lower degree of sophistication than the SPI. The goal of RAI, according to Rooy (1965), is to make it possible to compare precipitation deviations across regions. Using RAI and SPI, the current study was conducted to evaluate the frequency and intensity of drought for in the Harohar-Punpun Basin.

Is it not necessary to put materials and methods or meteorology for this section?

### **Study Area and Data Collection**

This study is focused on a region in India's Harohar-Punpun Basin that is prone to drought (Fig.1) and a description of the study area is shown in Table No. 1 (This is not a table it is a figure) or may be you can correct as Table 1 and Figure 1.

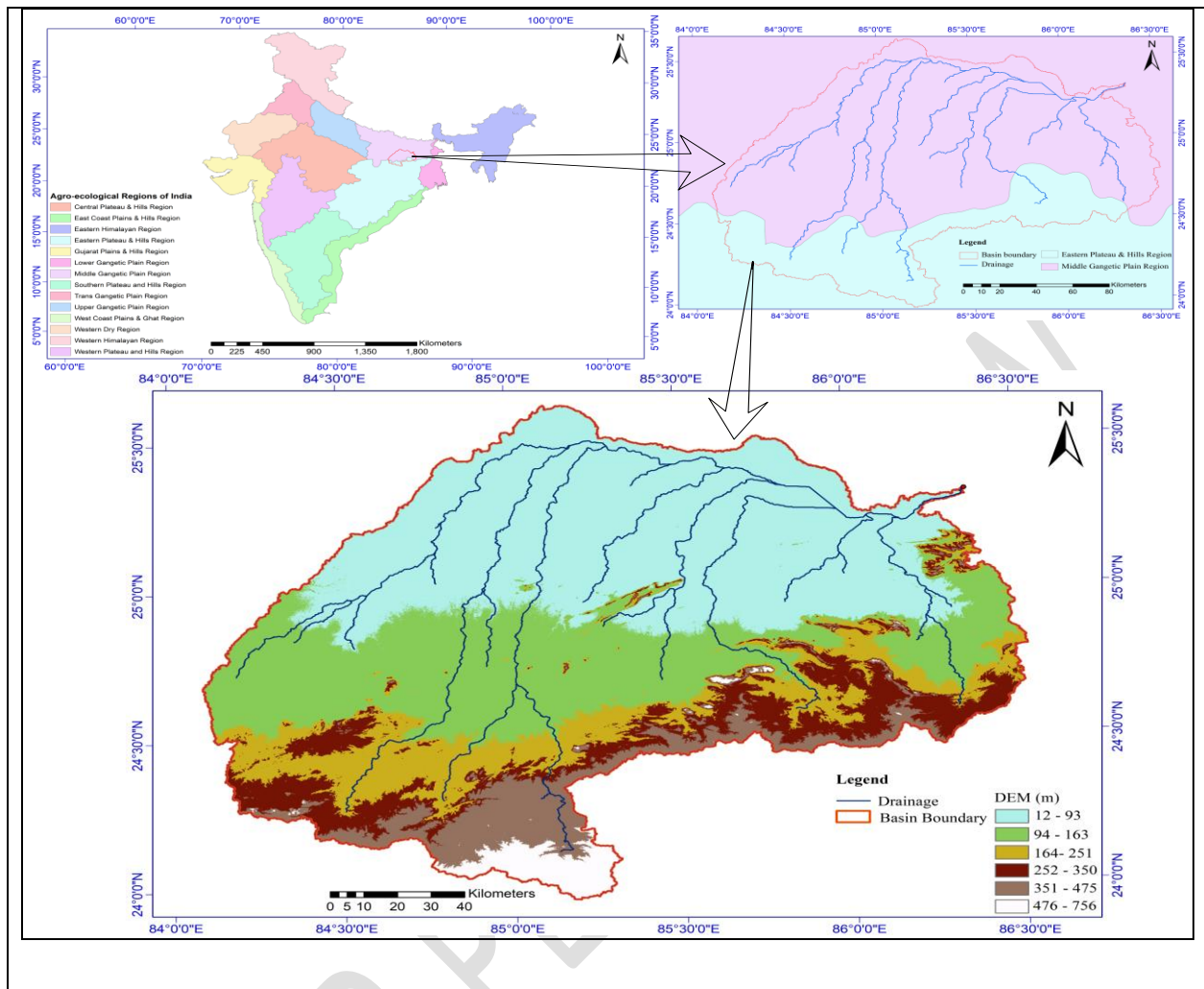


Fig 1: Study Area

Table No.1 Description of the study area

Characteristics	Punpun Basin	Harohar Basin
The Basin is situated between	Latitudes-24° 6' N & 25° 35' N Longitude-84° 0' E & 85° 19'E	Latitudes- 24°10' N & 25° 30' N Longitude- 84° 40'E & 86° 8' E
Geographical Area	9025.75sq.Km	14296.18 sq.Km
Covers the Area of Coverage	Patna district, Jehanabad district, Gaya district, Aurangabad district, Nalanda district, Palamu District, and Hazaribagh district.	Patna district, Nawada district, Jehanabad district, Gaya district, Munger District, Nalanda district, Hazaribagh district and Giridih district.

Rainfall (mm)	Average annual rainfall varies from 99cm near confluence with the Ganga (Patna districts) to 134 cm in the upper most reach (Palamu District).	Average annual rainfall varies from 99cm in the lower catchment of the river to 126cm in the hills of Hazaribagh district.
SoilType	Old alluvium grey, grayish-yellow, heavy textured cracking soil	Alluvium –calcareous, light-grey, medium grey to heavy textured soil.

IMD(India Meteorological Department) data of the mean annual and Daily rainfall data from 70 hydrological years (1951 - 2021) for different meteorological stations are provided.

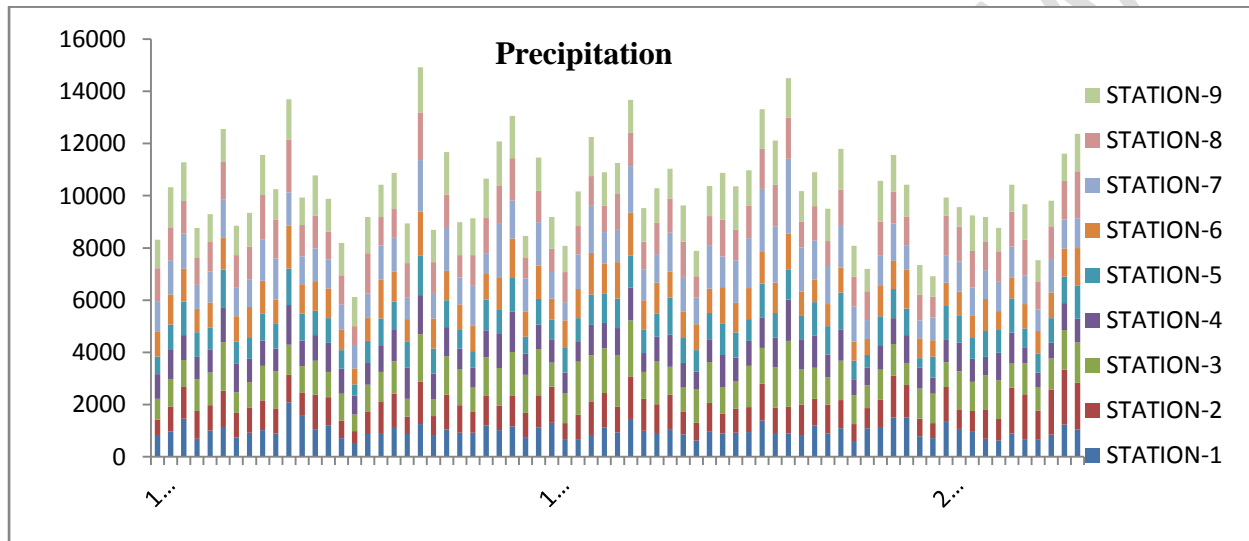


Fig. 2. Yearly precipitation (mm) from 1951 to 2021 for Punpun-Harohar Basin for different Stations.

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### Computation of the Standardized Precipitation Index (SPI)

McKee et al. (1993) developed the Standardized Precipitation Index (SPI) to monitor drought.

Thom (1966) found the gamma distribution to fit climatologically precipitation time series well.

The gamma distribution is defined by its frequency or probability density function:

$$G(x) = \frac{1}{\beta^\alpha \Gamma(\alpha)} x^{\alpha-1} e^{-\frac{x}{\beta}} \quad (1)$$

Where,  $\alpha > 0$ ,  $\alpha$  is a shape factor.  $\beta > 0$ ,  $\beta$  is a scale factor.

$$\Gamma(\alpha) = \int_0^\infty y^{\alpha-1} e^{-y} dy \quad (2)$$

Where,  $\Gamma(\alpha)$  is the gamma function.

Computation of the SPI involves fitting a gamma probability density function to a given frequency distribution of precipitation total for a station. From Thom (1966), the maximum likelihood solutions are used to optimally estimate  $\alpha$  and  $\beta$

$$\alpha = \frac{1}{4A} \left(1 + \sqrt{1 + \frac{4A}{3}}\right) \quad (3)$$

$$\beta = \frac{x}{\alpha} \quad (4)$$

$$A = \ln x - \frac{\sum \ln(x)}{n} \quad (5)$$

Where, n = number of precipitation observations.

The cumulative probability is given by:

$$G(x) = \frac{1}{\Gamma(\alpha)} \int_0^x t^{\alpha-1} e^{-t} dt \quad (6)$$

The gamma function is undefined for x=0 and a precipitation distribution may contain zeros, the cumulative probability becomes

$$H(x) = q + (1-q) G(x) \quad (7)$$

where q is the probability of a zero. If m is the number of zeros in a precipitation time series, Thom (1966) states that q can be estimated by m/n. The cumulative probability, H(x), is then transformed to the standard normal random variable Z with mean zero and variance of one, which is the value of the SPI. The classification shown in the following table is used to define drought intensities resulting from the SPI computation:

First tables and figures are cited in the body of a text followed by a caption of figures and tables

Table 2. Category of Standardized Precipitation Index (SPI) based on range values

SPI Values	Class
>2	Extremely Wet
1.5 to 1.99	Very Wet
1.0 to 1.49	Moderately Wet
-0.99 to 0.99	Near Normal
-1 to -1.49	Moderately Dry
-1.5 to -1.99	Severely Dry
< -2	Extremely Dry

### Computation of the Rainfall Anomaly Index (RAI)

The monthly RAI was calculated for specific years of the historical series aiming to analyze the distribution of rainfall in the years of greatest anomaly. RAI, developed and firstly used by Rooy (1965) and adapted by Freitas (2005), constitutes the following equations:

$$\text{RAI} = 3 \left[ \frac{N - \bar{N}}{\bar{M} - \bar{N}} \right] \text{ For positive anomalies.} \quad (8)$$

$$\text{RAI} = -3 \left[ \frac{N - \bar{N}}{\bar{x} - \bar{N}} \right] \text{ For negative anomalies.} \quad (9)$$

Where: N = current monthly/yearly rainfall, in order words, of the month/year when RAI will be generated (mm);

$\bar{N}$  = monthly/yearly average rainfall of the historical series (mm);

$\bar{M}$  = average of the ten highest monthly/yearly precipitations of the historical series (mm);

$\bar{x}$  = average of the ten lowest monthly/ yearly precipitations of the historical series (mm); and

positive anomalies have their values above average and negative anomalies have their values below average.

**Table 3. Category of Rainfall Anomaly Index (RAI) based on range values (where is in the text?)**

RAI range	Classification
Above 4	Extremely humid
2 to 4	Very humid
0 to 2	Humid
-2 to 0	Dry
-4 to -2	Very dry
Below -4	Extremely dry

Source: Freitas (2005) adapted by Araújo et al. (2009)

**What is the unit of RAI?**

## Results and Discussion

Where SPI-3 represents the drought index calculated over a time period of 3 months, SPI-6 over a time period of 6 months, and SPI-12 over a time period of 12 months, when the correlation matrices were examined, the strongest relationship was observed among the indices in the same time periods. As the time difference increases (monthly to yearly), the relationship between variables has weakened.

RAI and SPI values are obtained by normalization of rainfall data and give close results in both indices (Figure 3). However, RAI is simpler, as the calculation procedure does not need to be fitted to any theoretical distribution of data, according to SPI. Furthermore, it is more sensitive in detecting extreme drought and wetlands since it is fluctuating in a wider range.

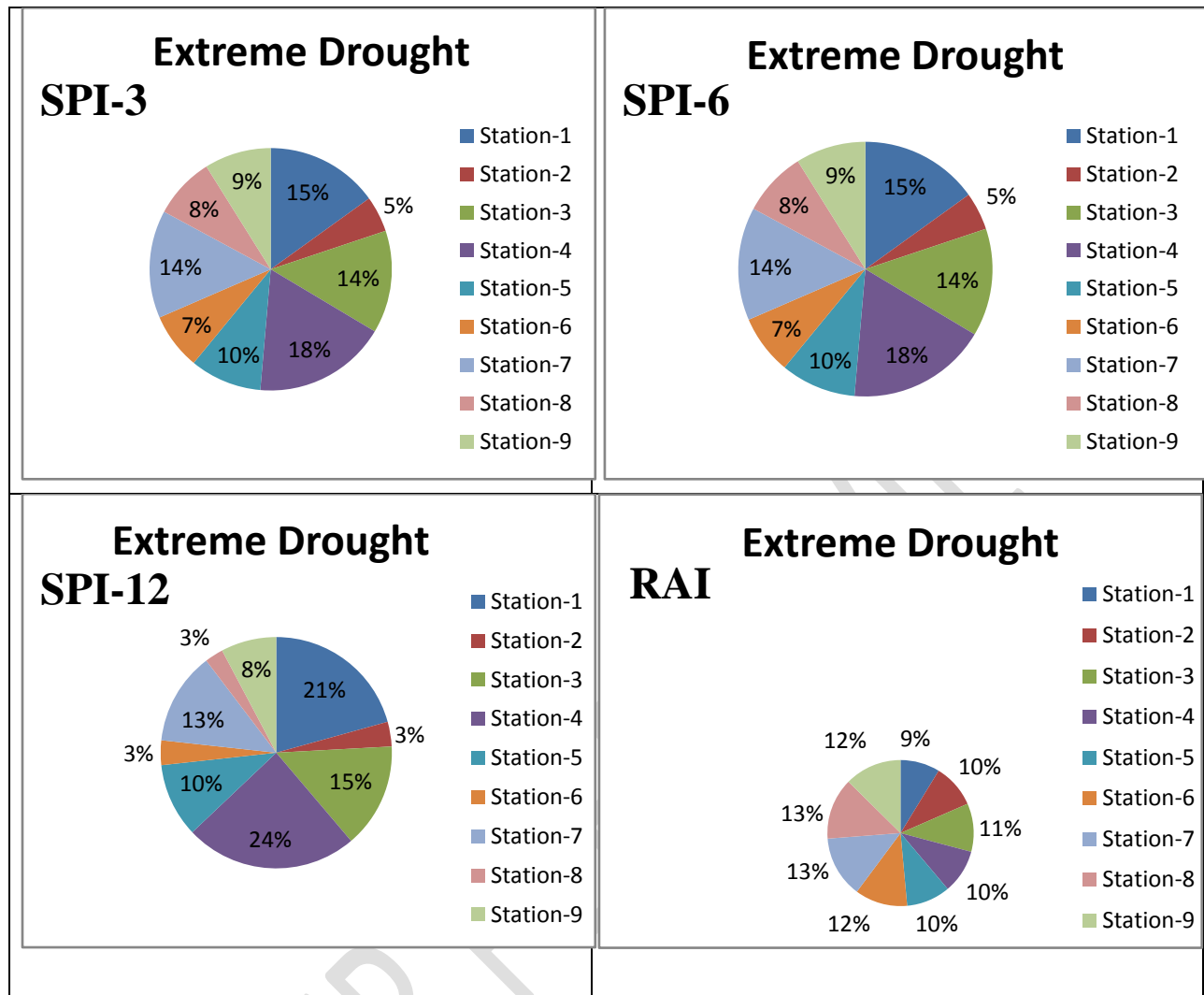
**5.1** Table No.4 The frequency of Extreme drought by RAI & SPI (3, 6, 12-Month)

<b>SPI-3</b>	Extreme	Frequency (%)
Station-1	14	1.65
Station-2	13	1.53
Station-3	11	1.30
Station-4	12	1.42
Station-5	15	1.77
Station-6	15	1.77
Station-7	6	0.71
Station-8	18	2.12
Station-9	18	2.12

<b>SPI-6</b>	Extreme	Frequency (%)
Station-1	22	2.61
Station-2	7	0.83
Station-3	20	2.38
Station-4	26	3.09
Station-5	14	1.66
Station-6	11	1.31
Station-7	21	2.49
Station-8	12	1.43
Station-9	13	1.54

<b>SPI-12</b>	Extreme	Frequency (%)
Station-1	24	2.89
Station-2	4	0.48
Station-3	17	2.05
Station-4	28	3.37
Station-5	12	1.45
Station-6	4	0.48
Station-7	15	1.81
Station-8	3	0.36
Station-9	9	1.08

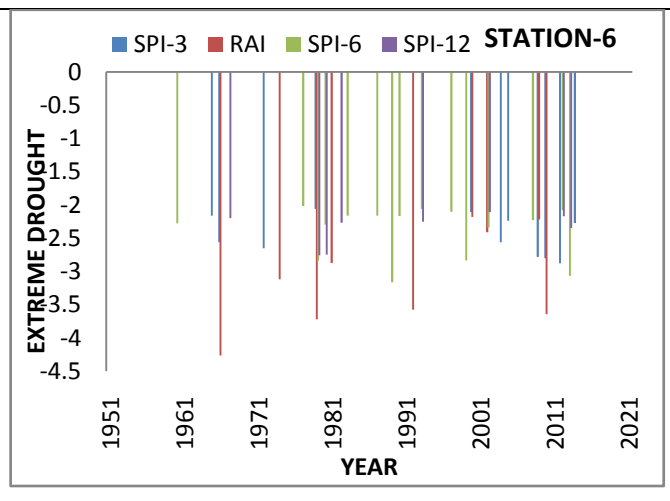
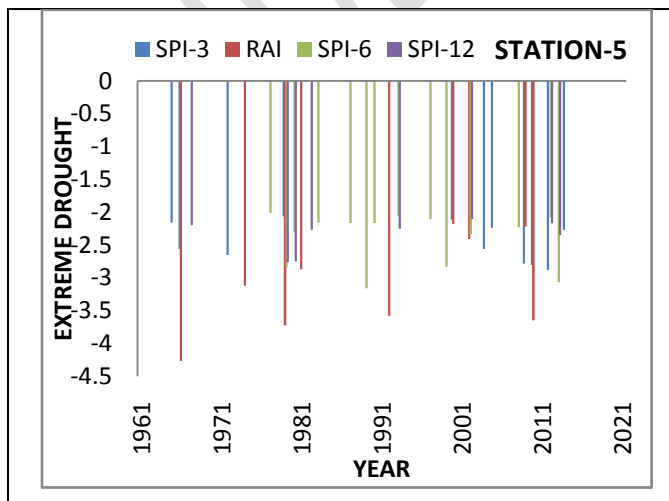
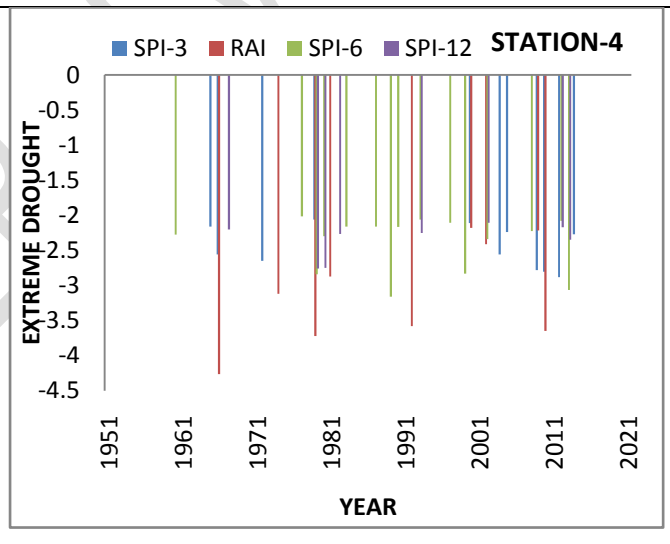
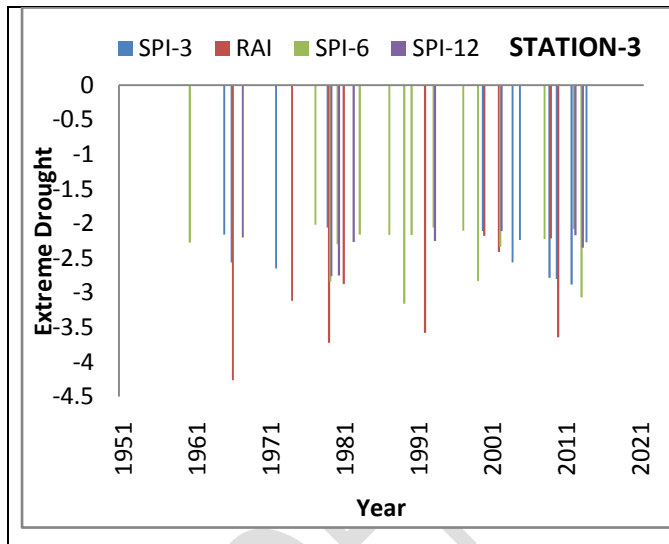
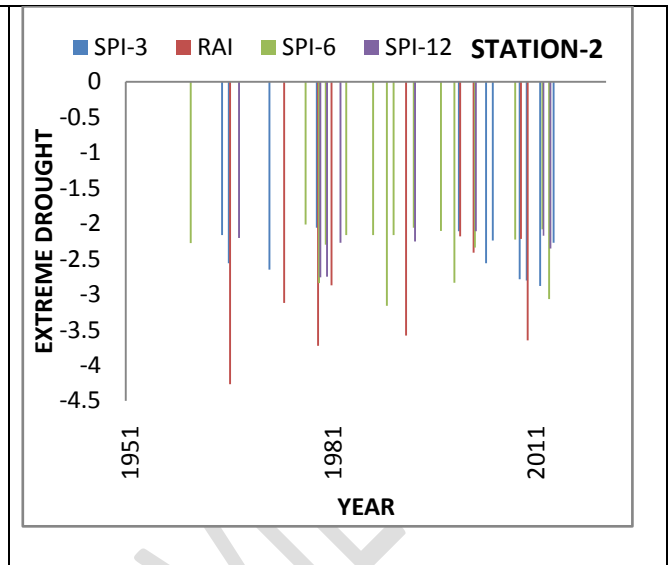
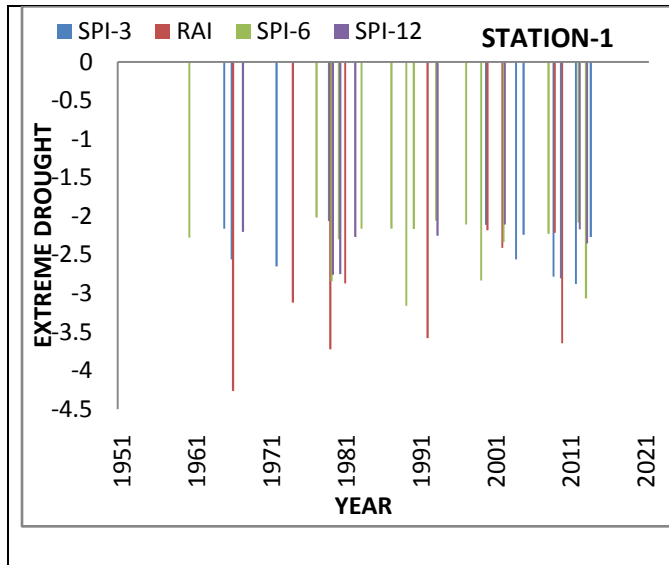
<b>RAI</b>	Extreme	Frequency (%)
Station-1	9	12.68
Station-2	10	14.08
Station-3	11	15.49
Station-4	10	14.08
Station-5	10	14.08
Station-6	12	16.90
Station-7	14	19.72
Station-8	14	19.72
Station-9	13	18.31

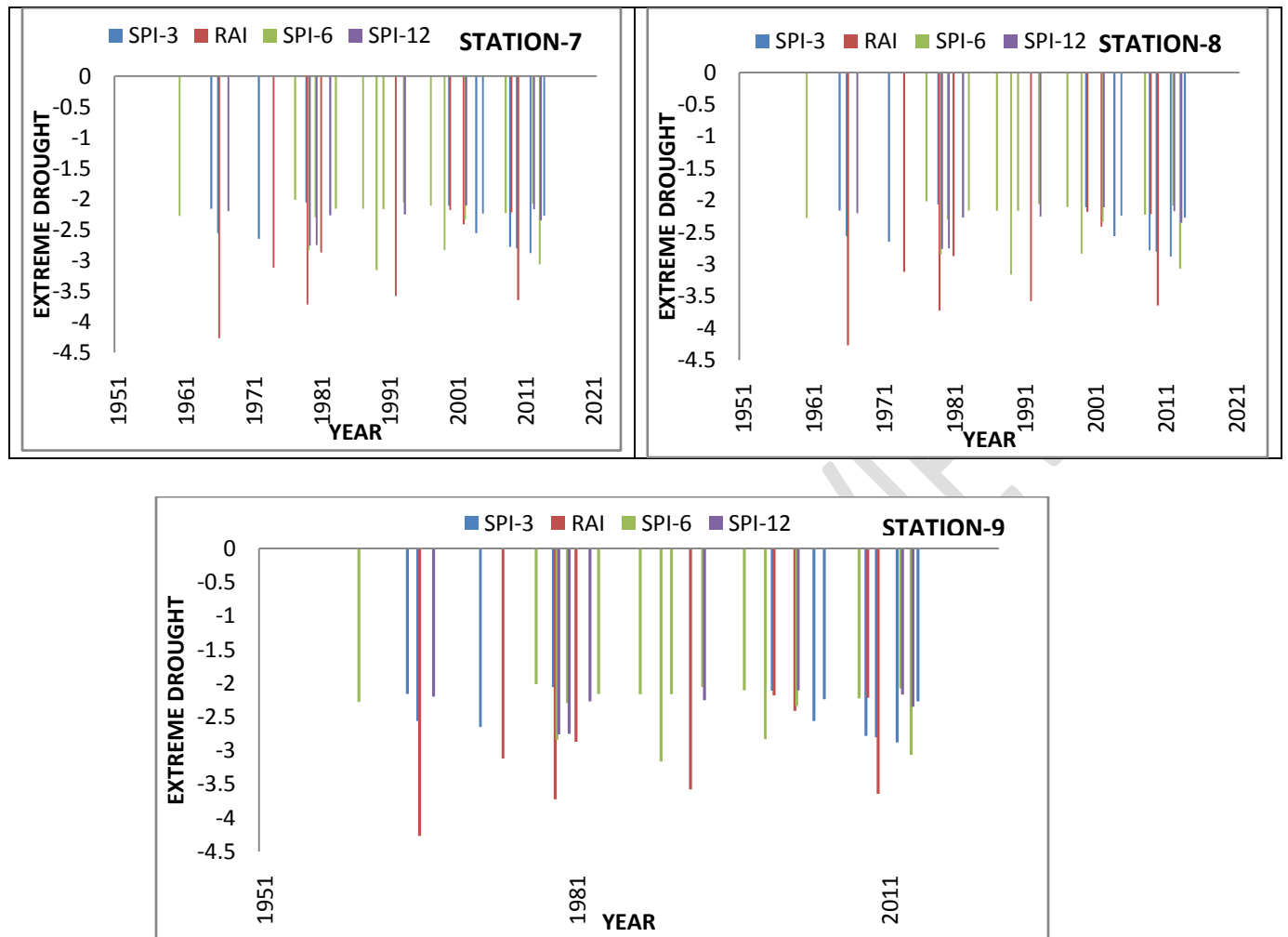


**Fig No. 3 Extreme Drought events detected by the RAI & SPI (3, 6, 12-Months)**

The same comment apply to this figure

The frequency of identified extreme drought events using RAI and SPI (3, 6, 12-Month) for various stations is depicted in Figure 3. In station-8 and station-9, the events were higher compared to the other stations, at 15%, while in station-5 and station-6; it was 12% for SPI-3. In SPI-6, station-1 exhibited a slightly higher value than station-7, with values of 2.69% and 2.49%, respectively. For SPI-12, station-4 and station-1 demonstrated higher values than the others, at 3.37% and 2.89%, respectively, during the period 1951-2021. In terms of RAI, the frequency was higher than that of SPI (3, 6, and 12). This represented a higher value of 19.72% for station-7 and station-8.





**Fig No4 - 3,6,12 - month SPI& RAI of extreme drought for Punpun-Harohar basin for Station-1 & Station-9**

### Comparison of SPIs and RAI

At Station-1, the RAI value outperformed SPI. The highest RAI value was -3.19 in 1972, while the lowest was -2.22 in 2009. Likewise, the values for SPI-3, SPI-6, and SPI-12 were -3.19 in 1972, -3.34 in 1979, and -2.89 in 1979, respectively. At Station-2, the RAI value was higher than the others, which were -4.15 in 2011. The values of SPI-3, SPI-6, and SPI-12 were -2.55, -2.63, and -2.50 in 1982, 1966, and 1967. At Station-3, SPI-3, SPI-6, SPI-12, and RAI values were similar but slightly higher for SPI-3, which was -3.78 in 2008. At Station-4, Station-5, Station-6, Station-7, Station-8, and Station-9, the RAI value was greater than that of all stations and SPIs. Which has the most negative anomalies concentrated in the southern portion of the study area: the values were -5.08, -4.06, -3.79, -4.26, -3.96, and -4.01 in 2009, 1966, 2009, 1966,

1966, and 1957. In 1972, the RAI value was similar to that of 1957. This is under extremely dry category. At Station-3, the values of SPI-3, SPI-6, and SPI-12 were higher than in other stations. These were -3.78, -3.91, and -3.03 in 2008, 2006, and 1967. The entire study area falls under the dry RAI and SPIs category.

### **Conclusion**

The current study aims to identify and assess the frequency of extreme drought in the southern part of the Harohar-Punpun Basin, Bihar. The RAI (Rainfall Anomaly Index) and SPIs (Standardized Precipitation Indices) methods were employed to evaluate extreme drought conditions in the study area. Analysis of rainfall data indicates that 1951 and 2021 experienced the lowest average rainfall, with recorded values of 990 mm and 1340 mm, respectively. Among the rainfall stations, stations 7, 2, and 3 exhibited the highest annual average rainfall, while stations 8, 9, and 6 demonstrated a medium annual average in the study area. The comprehensive investigation highlights the susceptibility of the southern part of the study area to prolonged dry spells. The RAI, serving as a metric for Rainfall Anomaly, facilitates comparisons between SPIs across different regions and allows for the assessment of drought frequency. Consequently, RAI proves to be a valuable tool in understanding the historical rainfall patterns of specific locations. The results demonstrate that the RAI value yielded more accurate outcomes compared to SPIs. This study holds significance in identifying potential drought-prone areas and serves as a valuable resource for planning and implementing mitigation strategies.

### **REFERENCE**

1. Bhunia, P., Das, P., Maiti, R., 2019. Meteorological Drought Study through SPI in Three Drought Prone Districts of West Bengal. India. *Earth Syst. Environ.* 4, 43–55. doi: 10.1007/s41748-019-00137-6.
2. Change, I. P. O. C., 2007. Fourth assessment report. IPCC, Ge.
3. Fernandes DS, Heinemann AB, Paz RL, Amorim AO, Cardoso AS. *Índices para a Quantificação da Seca*. Santo Antônio de Goiás, Embrapa Arroz e Feijão, 2009.
4. Freitas mas. Um sistema de suporte a decisão para o monitoramento de secas meteorológicas em regiões semiáridas. *Rev. Tecnol.* 2005; (suppl 19): p. 84-95.
5. Hao, Z. and Singh, V.P., 2018. Drought characterization from a multivariate perspective: A review. *Journal of Hydrol.*, 527, 668-678.

6. Hayes M J., 2006. Comparison of Major Drought Indices. National Drought Mitigation Center.<http://drought.unl.edu/Planning/Monitoring/ComparisonofIndicesIntro.aspx>.
7. Heim, R.R., 2002. A review of twentieth-century drought indices used in the United States. *Bull.Amer. Meteor. Soc.*, (83), 1149-1165.
8. Hou, Y.Y., He, Y.B., Liu, Q.H. and Tian, G.L., 2007. Research progress on drought indices. *Chin. J. Ecol.* 26, 892-897.
9. Keyantash J, Dracup JA (2002) The quantification of drought: an evaluation of drought indices. *Bull Am Meteorol Soc* 83:1167–1180.
10. Kim, T.W., Jehanzaib, M., 2020. Drought risk analysis, forecasting and assessment under climate change. *Water (Basel)* 12 (7), 12. doi: 10.3390/w12071862.
11. McKee, T.B., Doesken, N.J. and Kleist, J., 1993. The relationship of drought frequency and duration to time scales. *Proceedings of the 8th Conference on Applied Climatology*. AMS, Boston, MA, pp. 179-184.
12. Palmer, W. C., 1965. Meteorological Drought. U.S. Department of Commerce Weather Bureau Research Paper, 45, 1-58.
13. Palmer, W.C., 1968. Keeping track of crop moisture conditions, nationwide: the Crop Moisture Index. *Weatherwise* 21, 156-161.
14. Selvaraju, R., & Baas, S.E. (2007). *Climate variability and change: adaptation to drought in Bangladesh: a resource book and training guide*.
15. Smakhtin, V.U. and Schipper, E.L.F., 2008. Droughts: the impact of semantics and perceptions. *Water Policy*, 10, 131-143.
16. Sun, Y., Solomon, S., Dai, A., Portmann, R.W., 2006. How often does it rain? *J. Clim.* 19, 916–934. doi: 10.1175/JCLI3672.1.
17. Tirivarombo, S. and Hughes, D.A., 2011. Regional droughts and food security relationships in the Zambezi River Basin. *Phys. Chem. Earth*, 36, 977-983.
18. Van Rooy M (1965) A rainfall anomaly index independent of time and space. *Notes* 14:6
19. Vicente-Serrano S M, Begueria S, Lopez-Moreno J I. 2010. A multi-scalar drought index sensitive to global warming: The Standardized Precipitation Evapotranspiration Index. *Journal of Climate* 23: 1696-1718.
20. Zargar A, Sadiq R, Naser B, Khan F I. 2011. A review of drought indices. *Environmental Reviews* 19: 333-349.

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