

Actual Crop Evapotranspiration Estimation of Wheat Crop Using SEBAL Algorithm and Remotely Sensed Data .

Abstract

Estimation of evapotranspiration is imperative for effective forest, irrigation and water resources management as well as to increase yields and for better crop management. This study aims to evaluate the effectiveness of the Surface Energy Balance Algorithm for Land (SEBAL) in estimating actual evapotranspiration for 15 years of wheat crop in the 2A distributary of Mahanadi canal Command which is situated in Dhamtari district of Chhattisgarh state of India. The Landsat 7 and 8 satellite images from February – March (2007 -2021) were used to acquire the coefficients of the respective bands. To validate the outcomes from the SEBAL algorithm, FAO Penman-Monteith methods were employed to calculate the evapotranspiration values and evaluated using suitable performance metrics. The evaluation using SEBAL, along with the FAO-Penman–Monteith method, showed that SEBAL has a sufficient accuracy for estimating ET. The results showed that the SEBAL generated evapotranspiration values are in high agreement with the FAO Penman-Monteith method registering the highest correlation value ($R^2 = 0.826$) and the corresponding Root Mean Square Error was (RSME =0.4386), Normalized Root Mean Square, (NRMSE= 0.386), Mean Absolute Error (MAE=1.17) and Nash Sutcliffe Efficiency(ENS = 0.843) for Wheat crop. In 2010 maximum ET_c and in 2021 minimum ET_c of wheat crop in the study area. The outcomes show that since the performance of the SEBAL algorithm in estimating the actual evapotranspiration using Landsat 7 and 8 satellite images is acceptable, the SEBAL algorithm could be a very convenient method. Moreover, it could easily be assimilated into farming management systems and precision agriculture for better decision-making and higher yield.

Key words: Remote sensing, GIS, SEBAL Algorithm, Thermal band, FAO Penman-Monteith.

INTRODUCTION

Over the last few decades with the reduction of water resources, India is facing critical problems with industrial and agricultural growth being impacted. In agriculture, water requirements are associated with irrigation water use. Prediction of irrigation water demand involves computation of many water balance factors, and evapotranspiration (ET) is one of its major components. ET is the combined loss of water from the soil as well as plants and it is a crucial component of the hydrologic cycle. It is difficult to consider these two processes separately since they occur almost simultaneously at varied rates with high spatial variability (Allen et al. 1998b). As a result, evapotranspiration estimates are important for hydrology, irrigation, forest and water resources management. The evapotranspiration drives the soil water-energy balance which is largely used in general circulation models and climate modelling. Consequently, river water flow forecasting, crop yield forecasting, irrigation management systems, river/ lake water quality are all dependent on evapotranspiration levels (Ozturk and Apaydin 1998; Bastiaans sen et al. 2005; Razaji et al. 2020; Yamaç 2021). For this reason, it is essential to accurately estimate the water budget (Goyal and Harmsen 2013; Shabani et al. 2020). Better and accurate evapotranspiration estimates would allow for effective irrigation planning and optimal water usage for other agricultural purposes (Sattari et al. 2020a). The evapotranspiration rate depends on many factors such as temperature, solar radiation, humidity, wind and vegetation (Allen et al. 1998b). Essentially, the reference evapotranspiration (ET_0) is calculated using techniques such as FAO Penman and Hargreaves methods or measured directly using a lysimeter (Valayamkunnath et al. 2018). So far, various methods have been developed including direct/field measurements and empirical equations for estimating evapotranspiration. However, ET_0 consists of a complex and nonlinear structure requiring multiple parameters for estimation (Sattari et al. 2021). This nonlinear and multi-parameter nature makes estimation methods tedious and time-consuming (Sattari, Apaydin, and Shamshirband 2020b). The disadvantage of the conventional method is that it can only provide accurate evapotranspiration assessments of a

homogeneous region provided a meteorological gauge station is in the vicinity. Another draw back is that ET_0 cannot be extrapolated to different sites.

Consequently, surface observation networks have been developed, yet it is not possible to make meteorological measurements at all places covering large areas. The meteorological observations do not provide true spatial evapotranspiration of an area (Bastiaanssen et al. 1997; Liu et al. 2011; Antonopoulos and Antonopoulos 2017). On the other hand, remote sensing techniques allow for filling in the gap in providing the much-needed spatially observed data (Rawat et al. 2019). The advancements in remote sensing techniques in recent years together with the accessibility to satellite images have allowed for alternative and reliable methods for evapotranspiration estimations at regional scales (Mao and Wang 2017). Evapotranspiration models based on remote sensing provide relatively accurate estimations of evapotranspiration in large areas with minimal use of terrestrial data (Bastiaanssen and Chandrapala 2003).

Numerous models have been developed to estimate evapotranspiration using remote sensing methods. Out of all the proposed models for estimating evapotranspiration (Allen, Tasumi, and Trezza 2007), the Surface Energy Balance Algorithm for Land (SEBAL) model has proven to be the most widely used amongst researchers in over 30 countries. This model was developed by Bastiaanssen and improved by Allen (Bastiaanssen et al. 1998). The SEBAL model has proven to estimate evapotranspiration with better accuracy. Its registered accuracies of 85% at a farm-scale while more than 95% accuracy has been recorded on a regional scale (Seneviratne et al. 2006).

All these studies prove the efficiency and accuracy of the SEBAL method in estimating evapotranspiration across different geographical areas, climates, and land cover conditions. The SEBAL model has a lot of merits making it useful for such applications. It uses a linear relationship between the earth's surface temperature and near-surface vertical temperature gradients selected by cold and dry pixels (Bastiaanssen et al. 2005). More importantly, the SEBAL algorithm needs little information from the incoming terrestrial data to calculate the actual

evapotranspiration in each pixel and can further estimate evapotranspiration for large areas (Bashir et al. 2008) making it advantageous for real-life farm-based applications. Despite a number of studies being performed with SEBAL algorithms in evapotranspiration, improvements in accuracy are required for better results. In addition, the SEBAL method has a lower sensitivity to meteorological inputs and could easily be used in areas with fewer meteorological stations. According to studies done so far, it is necessary to evaluate and verify evapotranspiration estimates for each environmental condition and agricultural field since limited research has been carried out in the area of satellite estimates of evapotranspiration for crops.

Study Area

2A Distributary lies between $20^{\circ} 42'04''$ and $21^{\circ} '29''$ N latitudes and $81^{\circ} 47'50''$ and $81^{\circ} 85'30''$ E longitudes. The geographical area of the distributary 2A was 200.05 KM² with Cultivable Command Area (C.C.A) of 12461.58 ha . The major crop of the area is paddy in kharif season and in Rabi season mostly summer paddy and wheat, chickpea, pulses (mainly gram and millets) and oil seeds are grown.

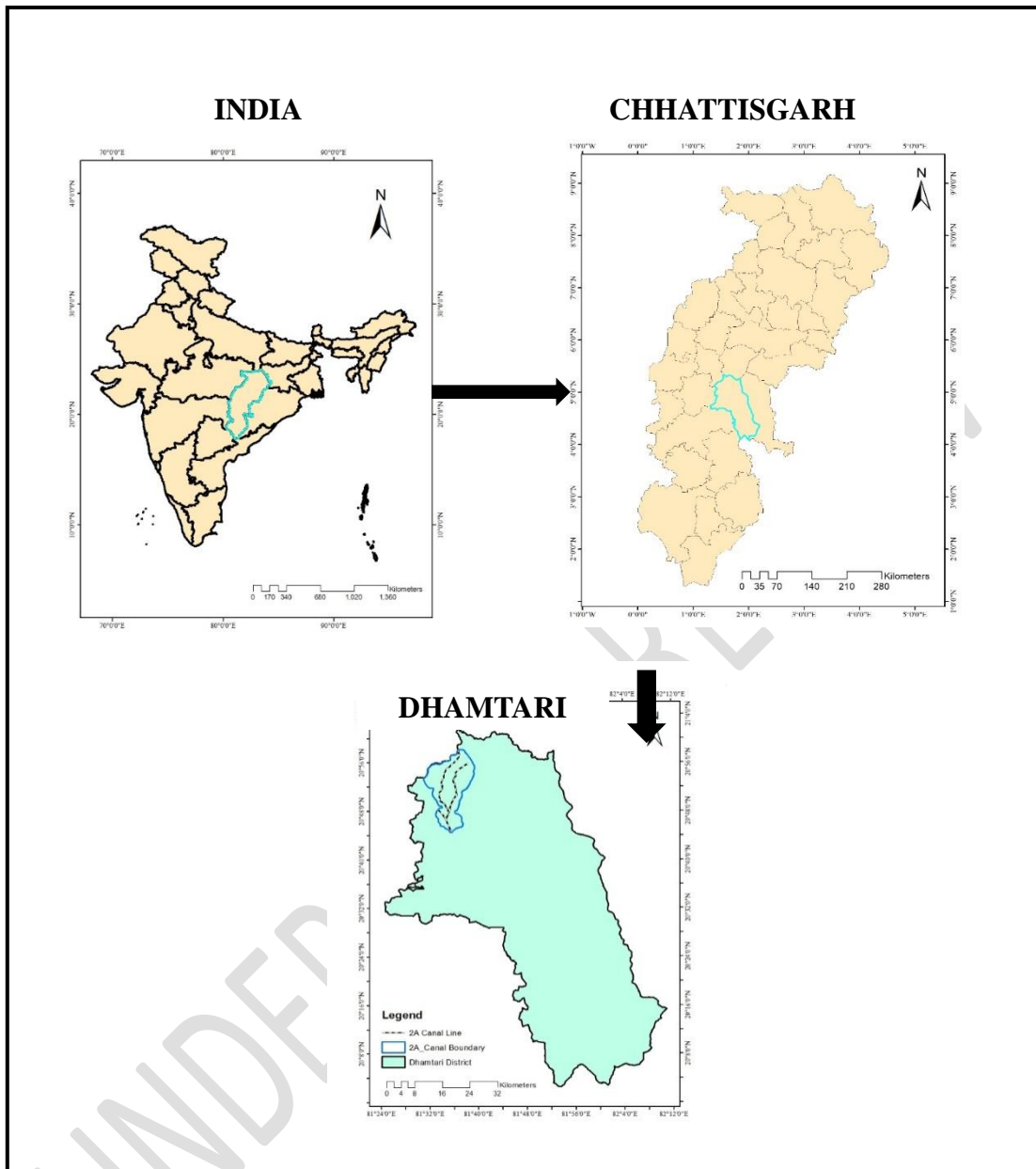


Fig.1. Location map of the study area

Materials and Methods

The Landsat 7 and 8 satellite images from February – March (2007 -2021) were used to acquire the coefficients of the respective bands. 15 years data were used for this study area. In this study , daily and hourly meteorological data from 2007-2021 were

used. The data included minimum and maximum humidity, average wind speed and sun hours.

Methodology

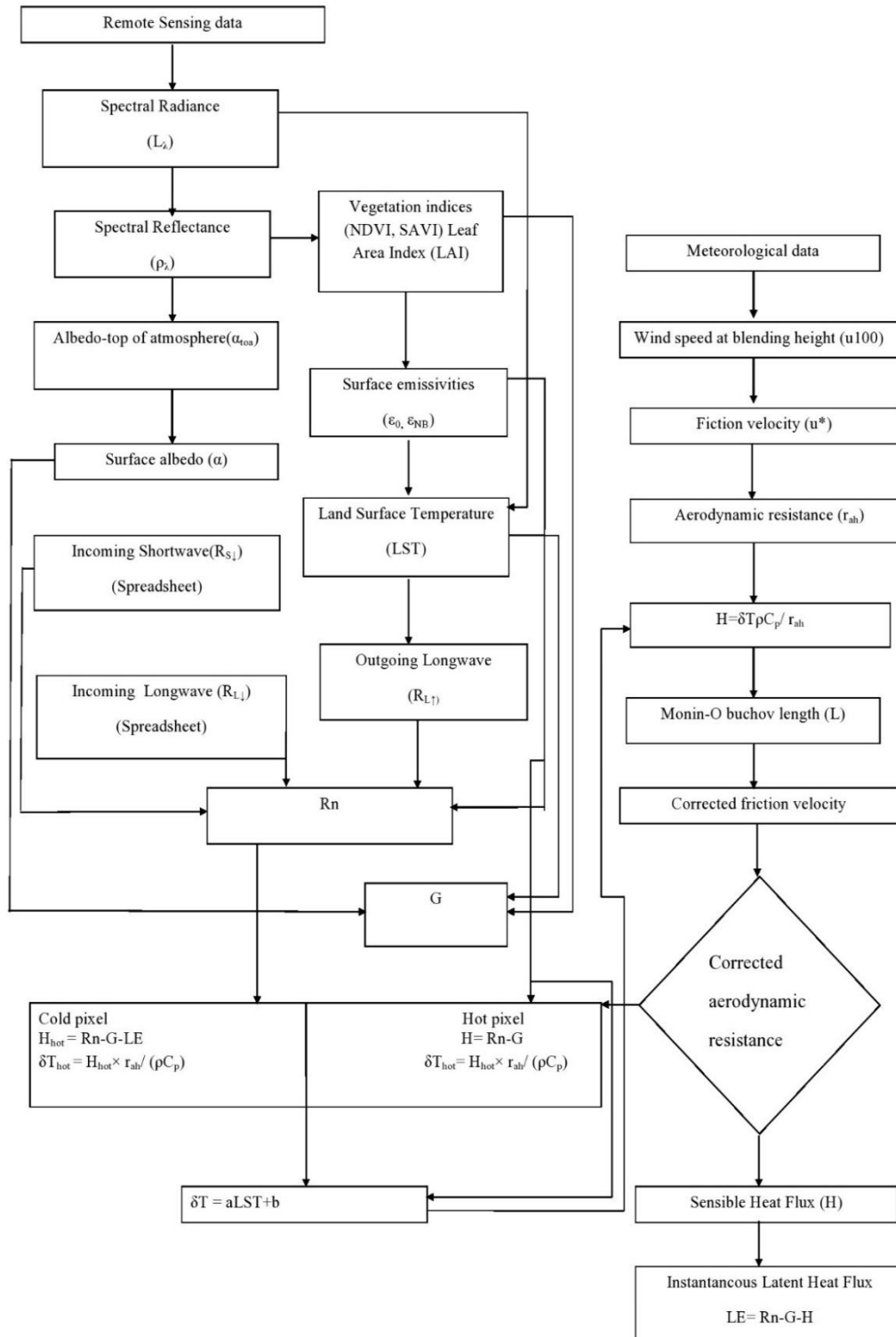


Fig.2. Flow chart of the computational steps of SEBAL algorithm (Bezerrz et al. 2015)

Surface energy budget equation

The total solar energy coming from the sun and atmosphere in the form of long and shortwave radiation is transformed and utilized for (a) soil heating, (b) surface environment heating (sensible heat flux to the atmosphere), and (c) transforming water into vapour (latent heat flux from the crop or soil surfaces). All the energy involved in the soil vegetation atmosphere interface can be given as the Energy Balance equation. The actual ET was obtained using the surface energy balance equation:

$$\lambda ET = R_n - G - H \quad \dots\dots(1)$$

Where, R_n is the net radiation at the surface (W/m^2), G is the soil heat flux (W/m^2), H is the sensible heat flux to the air (W/m^2), ET is the latent heat flux associated with ET (instantaneous value for the time of the satellite overpass, W/m^2).

Description of SEBAL Algorithm

The key input data for SEBAL consists of spectral radiance in the visible, near-infrared and thermal infrared part of the spectrum. In addition to satellite images, the SEBAL model requires routine weather data parameters (wind speed, humidity, solar radiation and air temperature).

Net Radiation (R_n)

The net radiation flux at the surface (R_n) represents the actual radiant energy available at the surface. It is computed by subtracting all outgoing radiant fluxes from all incoming radiant fluxes, which can be expressed as:-

$$R_n = (1-\alpha)R_{s\downarrow} + R_{L\downarrow} - R_{L\uparrow} - (1-\epsilon_0)R_{L\downarrow} \quad \dots\dots(2)$$

Where $RS\downarrow$ is the incoming shortwave radiation (W/m^2), α is the surface albedo (dimensionless), $R_{L\downarrow}$ is the incoming long wave radiation (W/m^2), $R_{L\uparrow}$ is the outgoing long wave radiation (W/m^2), and ϵ_0 is the surface thermal emissivity (dimensionless) (Water set al. 2002).

Surface Albedo (α)

Surface albedo (α) is defined as the ratio of the reflected radiation to the incident short wave radiation. On Surface it is a reflection coefficient. The albedo at the top of the atmosphere (α_{toa}) was computed. This is the albedo unadjusted for atmospheric transmissivity and is computed as follows:

$$\alpha_{toa} = \sum(\omega_{\lambda} \times \rho_{\lambda}) \quad \dots(3)$$

Where, ρ_{λ} is the computed reflectivity and ω_{λ} is a weighting coefficient for each band computed as follows:

$$\omega_{\lambda} = \frac{ESUN_{\lambda}}{\sum ESUN_{\lambda}} \quad \dots(4)$$

Where $ESUN$ is elevation of the sun. Albedo is defined as the ratio of the electromagnetic radiation reflected from the surface of the soil and the plant to the incident light emitted by the sun. Surface albedo is computed by correcting the α_{toa} for atmospheric transmissivity:

$$\alpha = \frac{\alpha_{toa} - \alpha_{path_radiance}}{\tau_{sw}^2} \quad \dots(5)$$

Values for $\alpha_{path_radiance}$ range between 0.025 and 0.04 and for SEBAL we recommend a value of 0.35 based on Bastiaanssen (1998). τ_{sw} includes the transmissivity of both direct solar beam radiation and diffuse (scattered) radiation to the surface. We calculate τ_{sw} assuming clear sky and

relatively dry conditions using an elevation-based relationship from FAO-56:

$$\tau_{sw} = 0.75 \times 2 \times 10^{-5} \times z \quad \dots(6)$$

Where, z is the elevation above sea level (m).

Incoming Shortwave Radiation (R_n)

Incoming shortwave radiation is the direct and diffuse solar radiation flux that actually reaches the earth's surface (W/m^2). Its value is computed as follows:

$$Rs_{\downarrow} = G_{sc} \times \cos \theta \times d_r \times \tau_{sw} \quad \dots(7)$$

Where

G_{sc} is the solar constant ($1367 W/m^2$),

$\cos \theta$ is the cosine of the solar incidence angle

d_r is the inverse squared relative earth-sun distance

τ_{sw} is the atmospheric transmissivity.

$$d_r = 1 + 0.033 \cos \left(\text{DOY} \frac{2\pi}{365} \right) \quad \dots(8)$$

d_r can be calculated as Solar incidence angle $\theta = 90 - \text{sun elevation angle}$ Atmospheric transmissivity

z is the elevation above sea level (m)

Outgoing Long wave Radiation (RL)

The outgoing long wave radiation is the thermal radiation flux emitted from the earth's surface to the atmosphere (W/m^2). It is computed in SEBAL through the following

Steps

Computation of vegetation indices of Normalized Difference Vegetation Index (NDVI), Soil Adjusted Vegetation Index (SAVI), and Leaf Area Index (LAI). The NDVI is the ratio of the differences in reflectivities for the near-infrared band (5) and the red band (4) to their sum for landsat-8 and for landsat-7 the near-infrared band (4) and the red band (3) to their sum:

$$\text{NDVI} = (R_1 - R_2) / (R_1 + R_2) \quad \dots(9)$$

The NDVI is a sensitive indicator of the amount and condition of green vegetation. Values for NDVI range between -1 and $+1$. Green surfaces have a NDVI between 0 and 1 and water and cloud are usually less than zero.

The SAVI is an index that attempts to "subtract" the effects of background soil from NDVI so that impacts of soil wetness are reduced in the index. It is computed as:

$$\text{SAVI} = (1 + L)(R_1 - R_2) / (R_1 + R_2) \quad \dots(10)$$

Where; L is a constant for SAVI. If L is zero, SAVI becomes equal to NDVI. A value of 0.5 frequently

appears in the literature for L . The LAI is the ratio of the total area of all leaves on a plant to the ground area represented by the plant. It is an indicator of biomass and

canopy resistance. LAI is computed for southern Idaho using the following empirical equation:

$$LAI = \frac{\ln\left(\frac{0.69 - SAVI}{0.59}\right)}{0.91} \quad \dots(11)$$

Where; $SAVI_{ID}$ is the SAVI calculated from Equation (11)

1. Computation of Surface emissivity (ϵ)

Surface emissivity (ϵ) is the ratio of the thermal energy radiated by the surface to the thermal energy radiated by a blackbody at the same temperature.

2. Computation of Outgoing Long wave Radiation ($R_{L\uparrow}$) This is computed using the Stefan-Boltzmann equation:

$$R_{L\uparrow} = \epsilon_0 \times \sigma \times T_s^4 \quad \dots(12)$$

Where ϵ_0 is the “broad-band” surface emissivity (dimensionless), σ is the Stefan-Boltzmann constant ($5.67 \times 10^{-8} \text{ W/m}^2/\text{K}^4$), and T_s is the surface temperature (K).

Choosing the “Hot” and “Cold” Pixels

The “cold” pixel is selected as a wet, well-irrigated crop surface having full ground cover by vegetation. The surface temperature and near-surface air temperature are assumed similar at this pixel. The “hot” pixel is selected as a dry, bare agricultural field where ET is assumed zero.

Incoming Long wave Radiation ($R_{L\uparrow}$)

The incoming long wave radiation is the downward thermal radiation flux from the atmosphere (W/m^2). It is computed using the Stefan-Boltzmann equation:

$$R_{L\uparrow} = \epsilon_0 \times \sigma \times T_s^4 \quad \dots(13)$$

Net surface radiation (R_n) is calculated using Equation (3.4)

Calculate Land Surface Temperature (LST)

For Landsat-7

This technique can only be used to process Landsat-7 (band 6) is used respectively.

Below given formula used to calculate :-

$$L_{\lambda} = \frac{(L_{MAX\lambda} - L_{MIN\lambda}) \times (Q_{CAL} - Q_{CAL\ MIN})}{(Q_{CAL\ MAX} - Q_{CAL\ MIN}) + L_{MIN\lambda}} \dots(14)$$

Where,

L_{λ} = Spectral Radiance (watts/meter)

Q_{CAL} = Quantized calibrated pixel value in DN

$L_{MAX\lambda}$ = Spectral Radiance scaled to $Q_{CAL\ MAX}$ (watts/meter)

$L_{MIN\lambda}$ = Spectral Radiance scaled to $Q_{CAL\ MIN}$ (watts/meter)

$Q_{CAL\ MAX}$ = Maximum Quantized calibrated pixel value (corresponding to $L_{MAX\lambda}$) in DN

$Q_{CAL\ MIN}$ = Minimum Quantized calibrated pixel value (corresponding to $L_{MIN\lambda}$) in DN

For Landsat-8

This technique can only be used to process Landsat-7 (band 6) and Landsat-8 data (band 10) is used respectively to estimate brightness temperature band (4-3) for Landsat-7 and band (5-4) for Landsat-8 are used to calculate NDVI.

Below given formula used to calculate :-

$$L_{\lambda} = \frac{(L_{max} - L_{min}) \times Q_{cal}}{(Q_{calmax} - Q_{calmin})} + L_{min} - Q_i \dots(15)$$

Where,

L_{max} is the maximum radiance ($Wm^{-2}sr^{-1}\mu m^{-1}$)

L_{min} is the minimum radiance ($Wm^{-2}sr^{-1}\mu m^{-1}$)

Q_{cal} is the DN value of pixel

$Q_{cal\ max}$ is the maximum DN value of pixels

$Q_{cal\ min}$ is the minimum DN value of pixels

O_i is the correction value for band 10

Soil Heat Flux (G)

Soil heat flux is the rate of heat storage into the soil and vegetation due to conduction. Estimates of G/R_n for agriculture surfaces is between 0.05–0.15.

$$\frac{G}{R_n} = \frac{T_s}{\alpha} \times [0.0032 \times \alpha + 0.0062 \times \alpha^2] \times [1 - 0.978 \times NDVI^4] \quad \dots(16)$$

Sensible Heat Flux (H)

Sensible heat flux is the rate of heat loss to the air by convection and conduction, due to a temperature difference. It is compute using the following equation for heat transport:

$$H = \frac{D \times C_D \times dt}{ra_h} = \rho \times C_p \times dt / ra_h \quad \dots(17)$$

Where

ρ is air density (kg/m^3),

C_p is air specific heat (1004 J/kg/K),

dT (K) is the temperature difference ($T_1 - T_2$) between two heights (z_1 and z_2),

ra_h is the aerodynamic resistance to heat transport (s/m).

Latent Heat Flux (λET), Instantaneous ET (ET_{inst}), and Reference ET Fraction (ET_{rF}) computation Latent heat flux is the rate of latent heat loss from the surface due to evapotranspiration. It can be computed for each pixel using the following Equation:

$$\lambda ET = R_n - G - H \quad \dots(18)$$

Where

λET is an instantaneous value for the time of the satellite overpass (W/m²).

An instantaneous value of ET in equivalent evaporation depth is computed as:

$$ET_{inst} = 3600 X \frac{\lambda ET}{\lambda} \quad \dots(19)$$

Where

ET_{inst} is the instantaneous ET (mm/hr), 3600 is the time conversion from seconds to hours,

λ is the latent heat of vaporization or the heat absorbed when a kilogram of water evaporates (J/kg) is computed as:

$$\lambda = [2.501 - 0.00236(TS - 273)] \times 10^6 \quad \dots(20)$$

Reference ET Fraction

The Reference ET Fraction (ET_rF) is defined as the ratio of the computed instantaneous ET (ET_{inst}) for each pixel to the reference ET (ET_r) computed from weather data:

$$ET_rF = \frac{ET_{inst}}{ET_r} \quad \dots(21)$$

Daily values of ET (ET_{24}) are often more useful than instantaneous ET.

Seasonal Evapotranspiration ($ET_{seasonal}$)

A seasonal evapotranspiration map was derived from the 24-hour evapotranspiration data by extrapolating the ET_{24} proportionally to the reference evapotranspiration (ET_r). ET_r is computed for a specific location and therefore does not represent the actual condition at each pixel. However, since ET_r is used only as an index of the relative change in weather, and therefore ET, for the image area. The ET_rF computed for the time of the image is constant for the entire period represented by the image. The length of the season for which ET is desired will be decided based on the period represented by each satellite image within the chosen season. The cumulative ET_r for the period represented by the image was computed by summing of daily ET_r values over the period as explained below :-

$$ET_{24} = ET_rF \times ET_{r-24} \quad \dots(22)$$

Where; ET_{r-24} is the cumulative 24-hour

ET_r for the day image. This is calculated by adding the hourly ET_r values over the day of the image (Waters et al. 2002).

Results and Discussion

Using SEBAL Algorithm or model various parameters value are found. The maximum value of Evapotranspiration was in 2010 and minimum in 2021 for wheat crop of 2A Distributary.

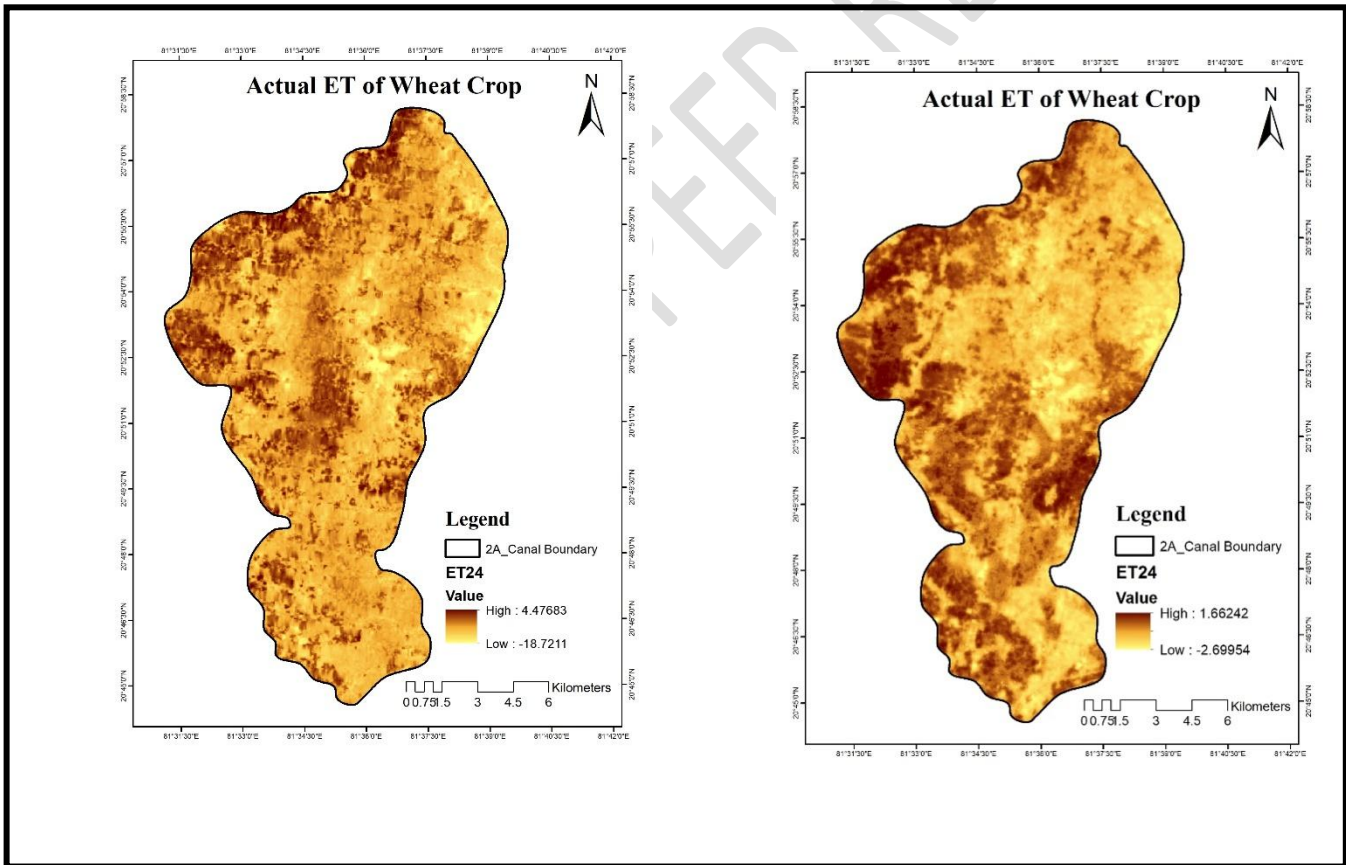
Table 1:-The results from SEBAL model for wheat crop of Rabi season

| Year | NDVI | Albedo % | T_s (°C) | R_N (W/m ²) | H (W/m ²) | G (W/m ²) | λET (W/m ²) | ET _{rF} | ET _{inst} | ET ₂₄ (mm/day) |
|------|-------|----------|------------|---------------------------|-----------------------|-----------------------|----------------------------------|------------------|--------------------|---------------------------|
| 2007 | 0.103 | 0.35 | 27.853 | 387.971 | 215.887 | 47.756 | 124.327 | 0.062 | 0.183 | 4.144 |
| 2008 | 0.105 | 0.35 | 28.676 | 353.083 | 222.261 | 41.104 | 89.7184 | 0.057 | 0.164 | 3.185 |
| 2009 | 0.075 | 0.35 | 33.994 | 511.684 | 258.749 | 65.176 | 187.759 | 0.053 | 0.221 | 3.033 |
| 2010 | 0.106 | 0.35 | 30.058 | 395.456 | 197.648 | 91.303 | 106.505 | 0.048 | 0.124 | 4.476 |
| 2011 | 0.127 | 0.35 | 27.118 | 211.711 | 123.142 | 51.897 | 36.671 | 0.014 | 0.042 | 3.893 |
| 2012 | 0.118 | 0.35 | 34.793 | 343.434 | 173.502 | 81.281 | 88.651 | 0.037 | 0.104 | 3.131 |
| 2013 | 0.139 | 0.35 | 29.487 | 227.975 | 144.824 | 28.942 | 54.191 | 0.021 | 0.063 | 4.117 |
| 2014 | 0.262 | 0.35 | 27.598 | 203.358 | 134.803 | 25.959 | 42.596 | 0.017 | 0.049 | 2.928 |
| 2015 | 0.265 | 0.35 | 26.568 | 178.612 | 124.071 | 25.614 | 28.927 | 0.011 | 0.033 | 4.171 |
| 2016 | 0.205 | 0.35 | 26.804 | 231.291 | 151.073 | 18.696 | 61.522 | 0.027 | 0.071 | 2.828 |
| 2017 | 0.216 | 0.35 | 25.885 | 162.818 | 119.491 | 22.464 | 20.863 | 0.084 | 0.024 | 3.935 |
| 2018 | 0.207 | 0.35 | 29.731 | 167.418 | 118.094 | 24.578 | 24.746 | 0.011 | 0.028 | 4.111 |
| 2019 | 0.215 | 0.35 | 25.571 | 191.945 | 123.528 | 15.921 | 52.496 | 0.024 | 0.061 | 2.605 |
| 2020 | 0.228 | 0.35 | 25.198 | 97.253 | 64.923 | 17.153 | 15.177 | 0.061 | 0.015 | 2.645 |
| 2021 | 0.155 | 0.35 | 23.806 | 119.097 | 85.616 | 11.653 | 21.828 | 0.011 | 0.025 | 0.645 |

As shown in Table 1, 15 images from 2007 to 2021 were used. The values of the parameters obtained by the SEBAL algorithm for Wheat Crop are presented here. Surface Albedo (α) is taken 0.35%, the Normalized Difference Vegetation Index (NDVI) range between -1 to +1 and all the values comes under range, (T_s) is Surface Temperature in degree Celsius maximum surface temperature is 34.793 (°C) in 2012 and minimum in 2021 it was 23.806 (°C), R_n is the net radiation flux at the surface (W/m²) maximum R_n is 511.684 (W/m²) in 2009 and minimum is 97.253 (W/m²) in

2020, (H) Sensible heat flux to the air (W/m^2) maximum H is 258.749 in 2009 and minimum is 64.923 in 2020, (G) is the Soil heat flux(W/m^2) maximum G is 91.303 in 2010 and minimum is 11.653 in 2021, λET is Instantaneous value for the time of the satellite overpass (W/m^2) maximum λET is 187.759 in 2009 and minimum is 15.177 in 2020, ET_rF is the Reference ET Fraction is maximum in 2017 (0.084) and minimum in 2015, 2018 (0.011), ET_{inst} is Instantaneous ET(mm/day) is maximum in 2009

(0.221) and minimum in 2020 (0.015), at the last ET_{24} (mm/day) are often more useful than instantaneous ET maximum in 2010 (4.476) and minimum in 2021 (0.645), all these value is obtained for the study area based on the studied methods. Map 1 of the maximum and minimum Evapotranspiration are shown.



Conclusion

Increasing population and existing climate change scenario is posing a major challenge to the global fresh water resource. This challenge is more visible in agriculture sector, especially of water stressed countries, as it is often the biggest user of fresh water supplies. Over the last few decades with the reduction of the water resources, India is facing critical problems with industrial and agricultural growth being impacted. Prediction of irrigation water demand involves computation of many water balance factors and evapotranspiration (ET) is one of its major components. ET is the combined loss of water from the soil as well as plants and it is a crucial component of the hydrologic cycle. The accurate estimation of the crop evapotranspiration (ET_c) is vital in the management and development of water resources. The scarcity of fresh water resources or the misuse of water in agriculture has made the efficient and profitable use of water even more important.

The present study was carried out in Dhamtari district, Kurud block of Mahanadi Command area in Chhattisgarh State of Northern part of Dhamtari district. The area lies between $20^{\circ}.74'04''$ and $21^{\circ}.08'29''$ N latitudes and $81^{\circ}.47'50''$ and $81^{\circ}.85'30''$ E longitudes. To validate the outcomes from the SEBAL algorithm, FAO Penman-Monteith methods were employed to calculate the evapotranspiration values and evaluated using suitable performance metrics. The evaluation using SEBAL, along with the FAO-Penman-Monteith method, showed that SEBAL has a sufficient accuracy for estimating ET. The results showed that the SEBAL generated evapotranspiration values are in high agreement with the FAO Penman-Monteith method registering the highest correlation value ($R^2 = 0.826$) and the corresponding Root Mean Square Error was (RSME =0.4386), Normalized Root Mean Square, (NRMSE= 0.386), Mean Absolute Error (MAE=1.17) and Nash Sutcliffe Efficiency (ENS = 0.843) for Wheat crop. In 2010 maximum ET_c and in 2021 minimum ET_c of wheat crop in the study area. We show that the results of this study can be applied in the studies of water resources management and appropriate irrigation management on farm level.

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